

Chemical Weathering and Soils

Fresh rocks and minerals that once occupied the outermost position reached their present condition of decay through a complex of interacting physical, chemical, and biological processes, collectively called *weathering*.



Chemical Weathering and Soils

1. new minerals created by the weathering processes,
2. minerals that resisted destruction, and
3. organic debris added to the weathering zone.



Chemical Weathering and Soils

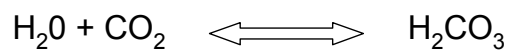
Functions of weathering

- gives rock lower strength and greater permeability,
- produces minor landforms,
- releases minerals in solution,
- Is first step in soil formation.

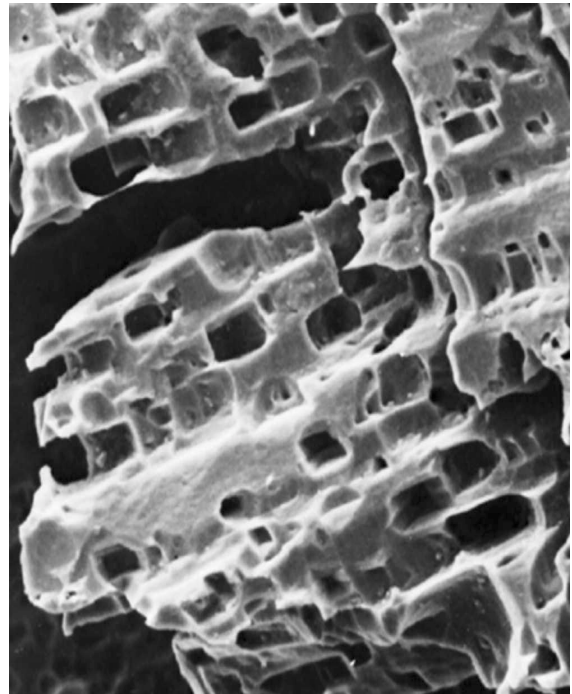
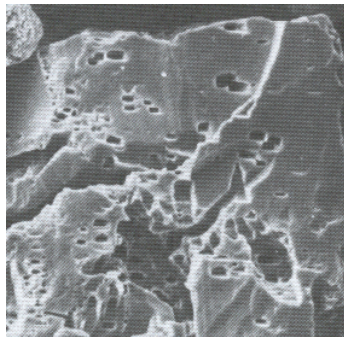
Chemical Weathering and Soils

Decomposition - group of processes that attempt to create substances that are more nearly stable in that environment.

- Driving force for decomposition is water.



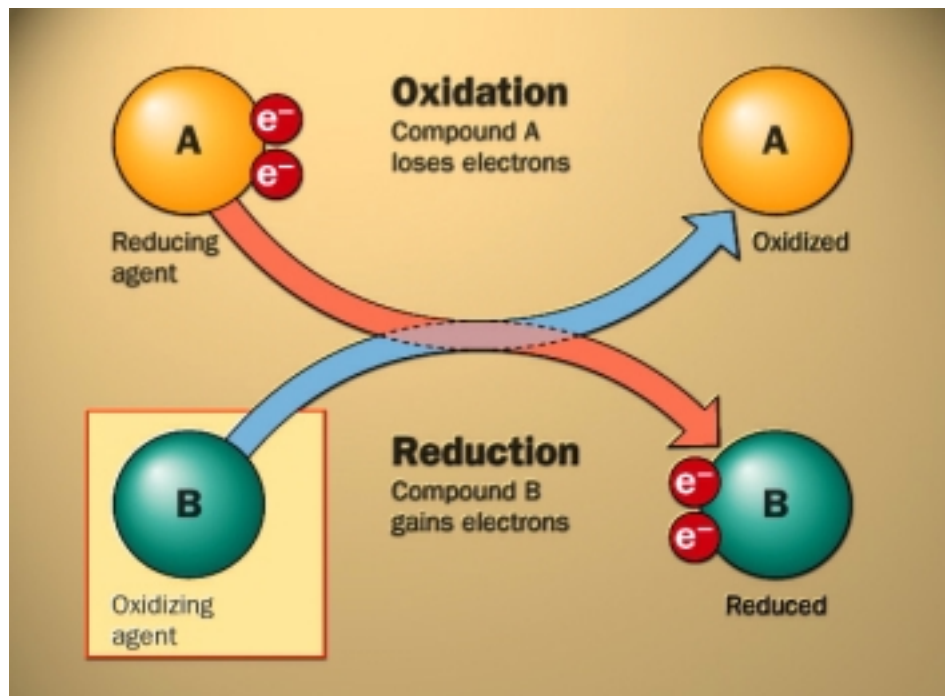
- How materials weather is related to how easily ions are released.
- sites of excess energy



Chemical Weathering and Soils

Processes of Decomposition

Oxidation and Reduction



Chemical Weathering and Soils

Processes of Decomposition

Oxidation and Reduction

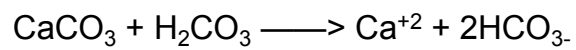
- iron is the most commonly oxidized mineral element
 Fe^{+2} (ferrous iron) \longrightarrow
 Fe^{+3} (ferric iron) or $2\text{FeO} + \text{O}_2 \longrightarrow \text{Fe}_2\text{O}_3$



Chemical Weathering and Soils

Processes of Decomposition

Solution - when atoms are dissolved from a mineral, the structure becomes unstable.

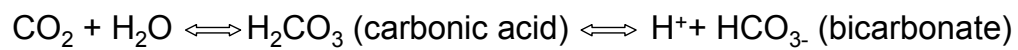


- bicarbonate represents the largest constituent of the dissolved load of most rivers

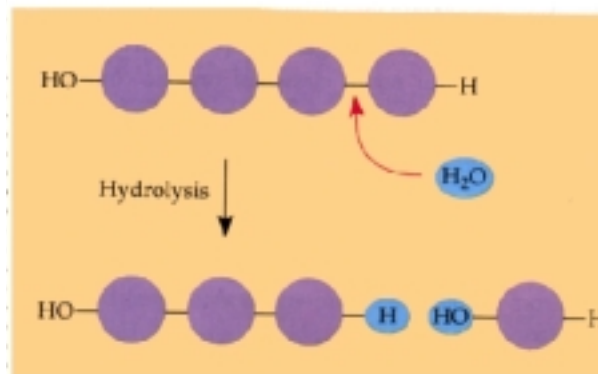
Chemical Weathering and Soils

Processes of Decomposition

Hydrolysis - a reaction between a salt and water to produce an acid and a base.



Hydrolysis

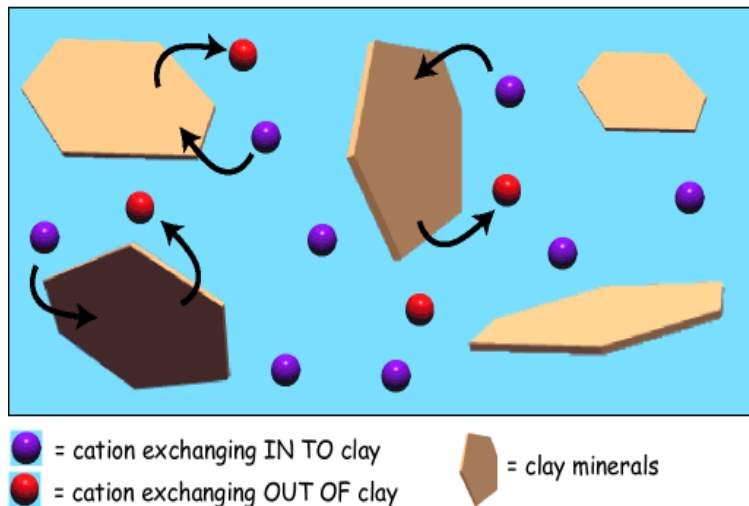


Chemical Weathering and Soils

Ion exchange - is the substitution of ions in solution for those held by mineral grains - most effective in clay minerals.

cation exchange capacity (c.e.c.) - exchange is governed by the composition and pH of the interstitial water as well as the type of ion in the exchangeable position.

Cation Exchange



Chemical Weathering and Soils

Representative Cation Exchange Capacities for Various Materials

<i>Material</i>	<i>Approximate cation exchange capacity (me/100 g dry weight)</i>
Organic matter	150–500
Kaolinite	3–15
Halloysite	5–10
Hydrated halloysite	40–50
Illite	10–40
Chlorite	10–40
Smectite	80–150
Vermiculite	100–150+
Palygorskite	5–30
Sepiolite	20–45
Allophane	25–70
Hydrous oxides of aluminum and iron	4
Feldspars	1–2
Quartz	1–2
Basalt	1–3
Zeolites	230–620

(Taken from Carroll¹⁵ and Grim³²; palygorskite and sepiolite data from Weaver and Pollard.⁷¹)

Chemical Weathering and Soils

Mobility of decomposed materials

Leaching

- removes in solution the constituents that have been separated from minerals
- affects the pH of fluids surrounding minerals
- transfer mechanism

Chemical Weathering and Soils

Mobility of decomposed materials

Fixation and retardation - mobility is considerably lower than might be expected.

Chelation - reaction between a metal ion and a complexing agent

- metallic ions that are extremely immobile under normal conditions can be mobilized by reacting with complexing agents

Chemical Weathering and Soils

Degree and rate of Decomposition

- End products
- Mineral stability
- Secondary minerals
- Estimates based on chemical analyses

Mean Lifetime of a 1 mm Crystal at 25°C and pH 5
(From Lasaga et al., 1994, Table 1 and References
Therein).

<i>Mineral</i>	<i>Lifetime (years)</i>
Quartz	34,000,000
Kaolinite	6,000,000
Muscovite	2,600,000
Epidote	923,000
Microcline	921,000
Prehnite	579,000
Albite	575,000
Sanidine	291,000
Gibbsite	276,000
Enstatite	10,100
Diopside	6,800
Forsterite	2,300
Jepheline	211
Anorthite	112
Wollastonite	79

Chemical Weathering and Soils

Soils

The vertical arrangement of the layers constitutes a diagnostic property of soils known as the ***soil profile***.



Relationship
between bedrock
and soil profile

Soil Profile

Bedrock



Chemical Weathering and Soils

Fundamental concepts of soil genesis (pedogenesis)

- Present day processes operate through space and time,
- Distinct regimes of soil processes produce distinct soils,
- Soil and its vegetation cover modify processes of land degradation,
- Clay is produced within the soil,
- Organo-mineral complexes are produced in the soil,

Chemical Weathering and Soils

Fundamental concepts of soil genesis (pedogenesis)

- In the course of pedogenesis, soil succession occurs,
- Complexity is more common than simplicity
- Little of the soil continuum is older than Tertiary and most no older than Pleistocene times,
- A knowledge of the Pleistocene is pre-requisite to understanding soils.

Chemical Weathering and Soils

The soil profile



O horizon

A horizon

E horizon (eluvial)

B horizon (illuvial)

C horizon

Bedrock

Chemical Weathering and Soils

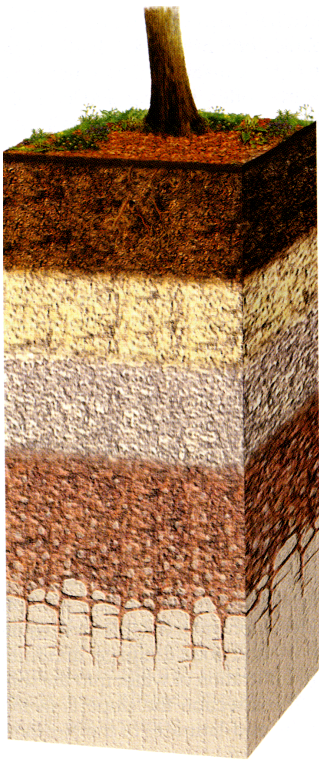
The soil profile



O Horizons – layers dominated by organic material.

Chemical Weathering and Soils

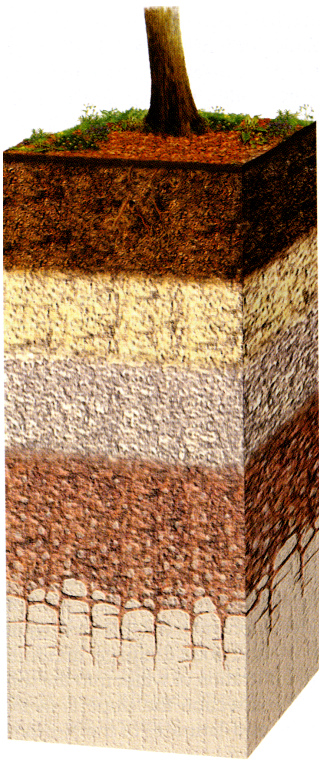
The soil profile



A horizons – Mineral horizons that formed at the surface or below an O horizon, that exhibit obliteration of all or much of the original rock structure and show one of more of the following:

Chemical Weathering and Soils

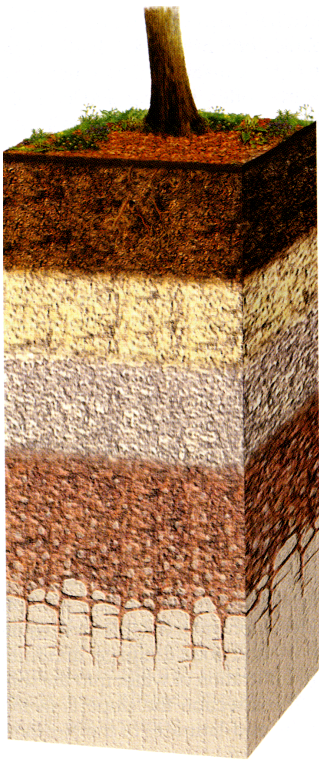
The soil profile



E horizon – underlies the O or A horizon and is characterized by intense leaching that removes Fe^{3+} or organic coatings from the mineral particles (*eluvial*).

Chemical Weathering and Soils

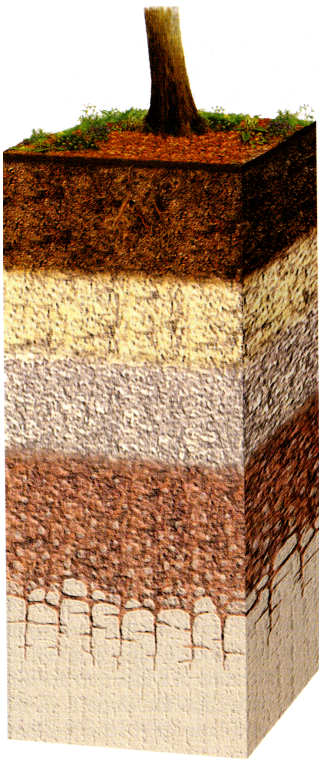
The soil profile



B horizons –generally show *illuvial* concentration of silicate clay, iron, aluminum, humus, carbonates, gypsum, or silica alone or in combination;

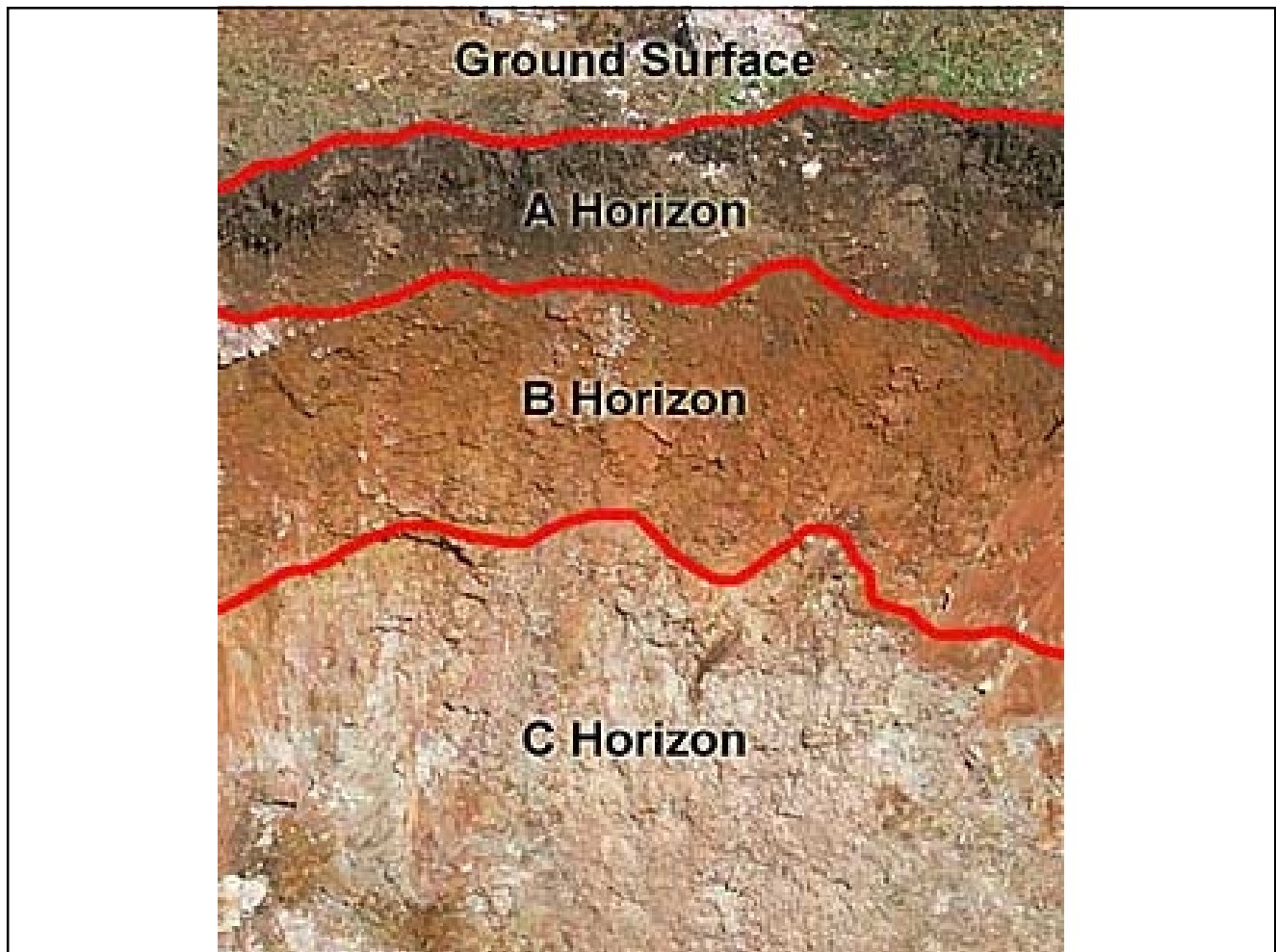
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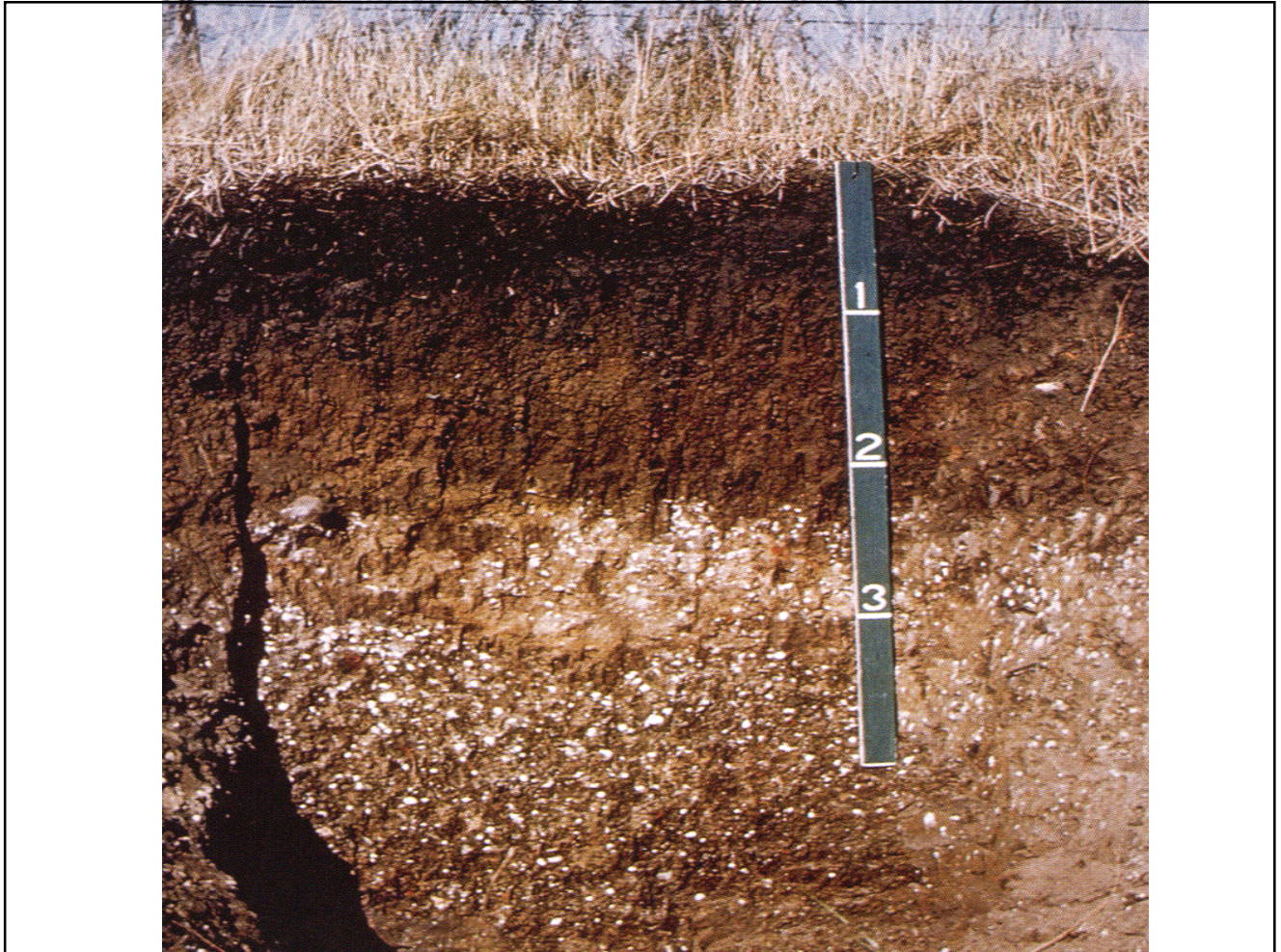
The soil profile



C horizon has no properties typical of the overlying horizons.

The **R horizon** is simply consolidated bedrock beneath the soil.









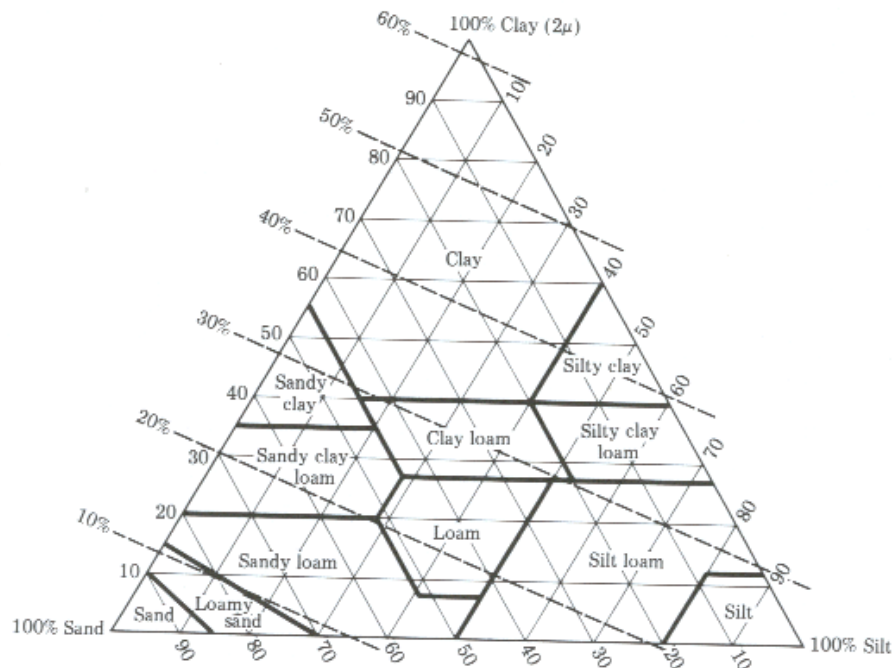




Chemical Weathering and Soils

Soil Properties

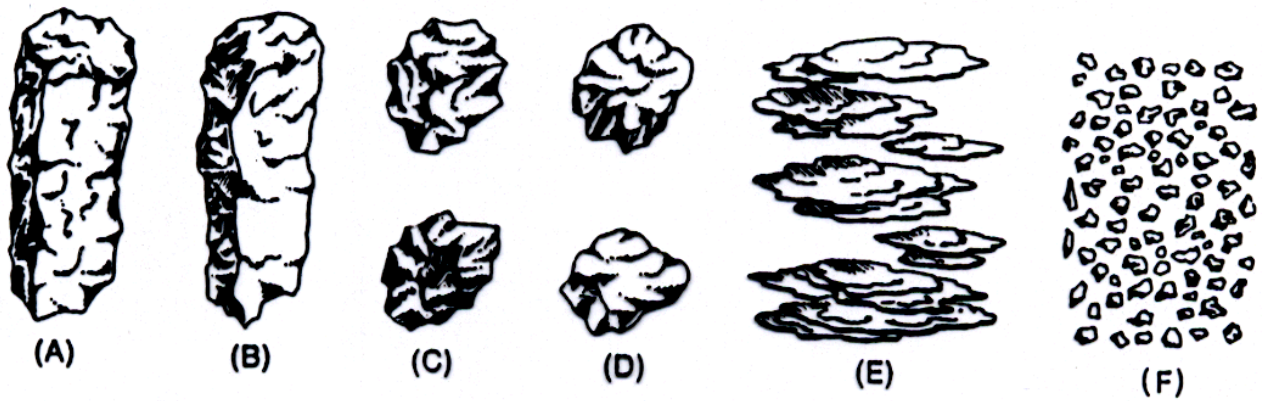
Texture - proportions of different particle sizes in a soil horizon.



Chemical Weathering and Soils

Soil Properties

Structure is a unique characteristic in that it designates the shape developed when individual particles cluster together into aggregates called *peds*.

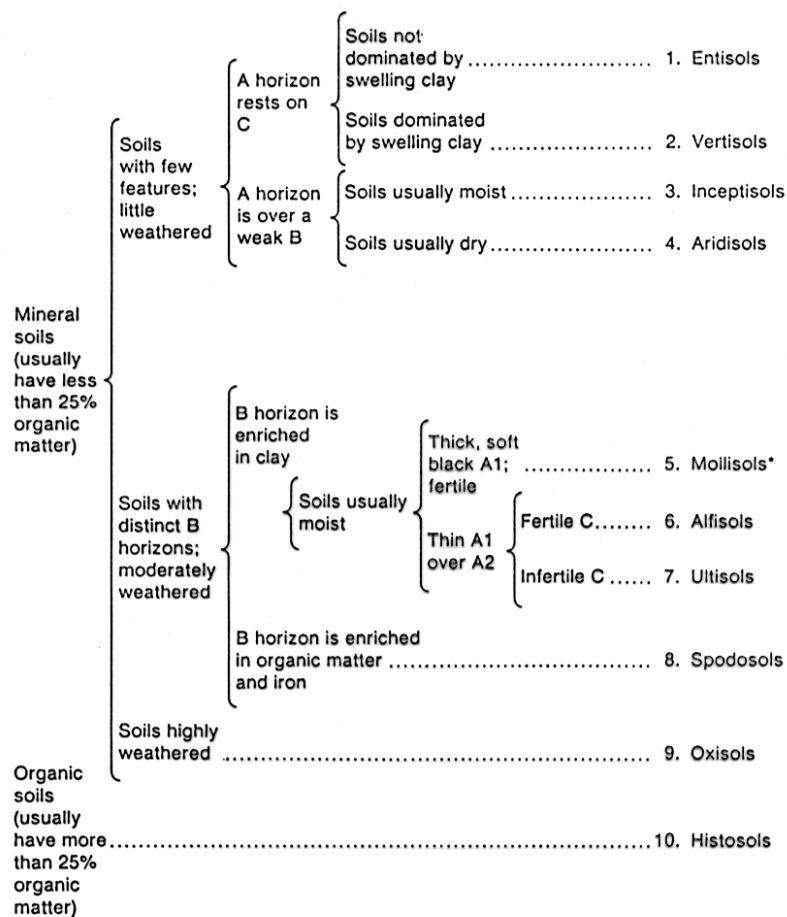


Chemical Weathering and Soils

Soil classification

major orders are:

Entisols	-	recent
Vertisols	-	inverts
Inceptisols	-	inception
Aridisols	-	arid
Mollisols	-	mollify (to lessen in intensity, to soften
Spodosol	-	podzol, odd
Alfisol	-	pedalfer
Ultisol	-	ultimate
Oxisol	-	oxide
Histosol	-	histology (tissue)



*Some Mollisols have a weak B horizon; not enriched by clay.

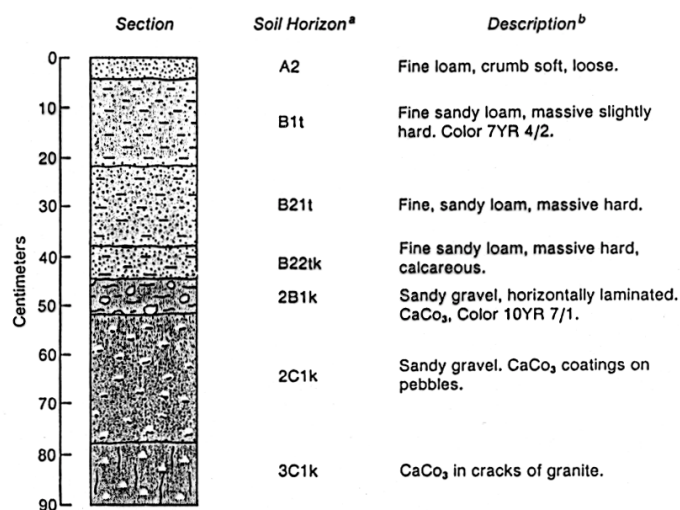
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Symbol ^a	Meaning
b	buried soil horizon
g	strong gleying
h	illuvial humus
ir	illuvial iron
k ^b	accumulation of alkaline earth carbonates, commonly CaCO ₃
m	strong cementation
p	plowing
t	illuvial clay
x	fragipan character

From Soil Survey Staff, 1975, 1981.

^aSymbols used with other profile designations (e.g., B2t, B1h, Cca).

^bFormerly designated as ca, a term many publications still use.

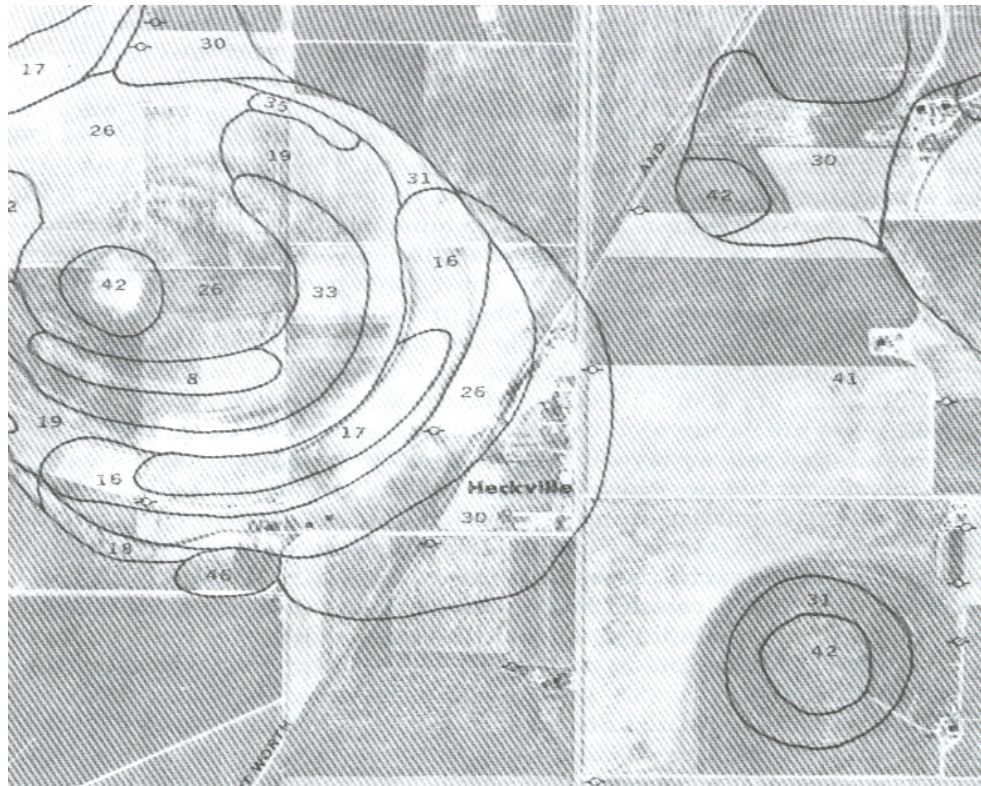


^a Lowercase k may now be used in place of ca.

^b Color from Munsell color charts (see Soil Survey Staff, 1975, pp. 463–69 for discussion).

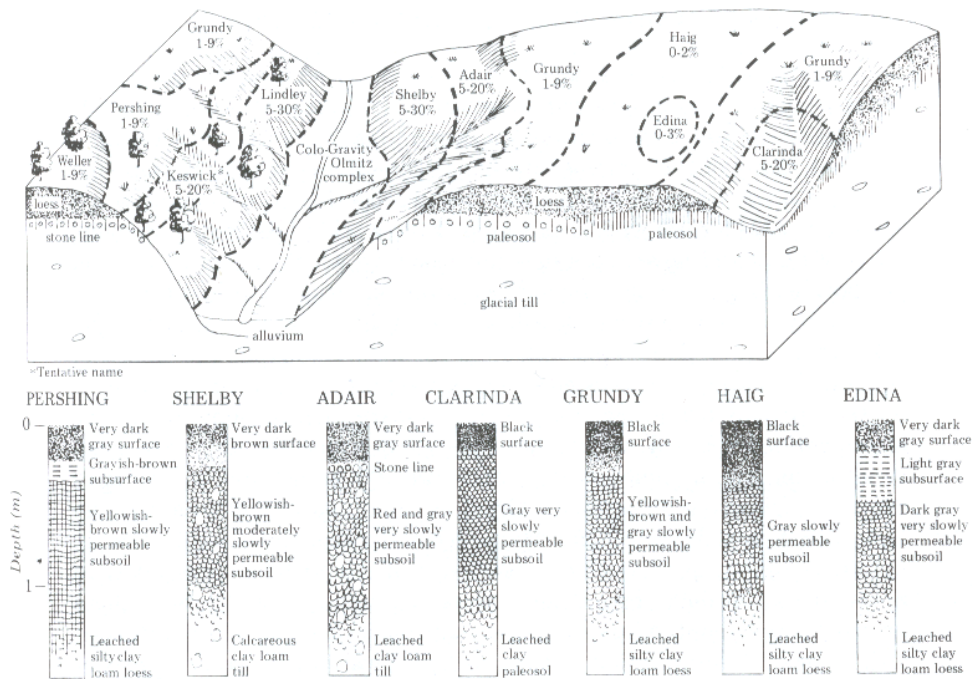
Chemical Weathering and Soils

How are soils mapped?



Chemical Weathering and Soils

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Chemical Weathering and Soils

How are soils mapped?

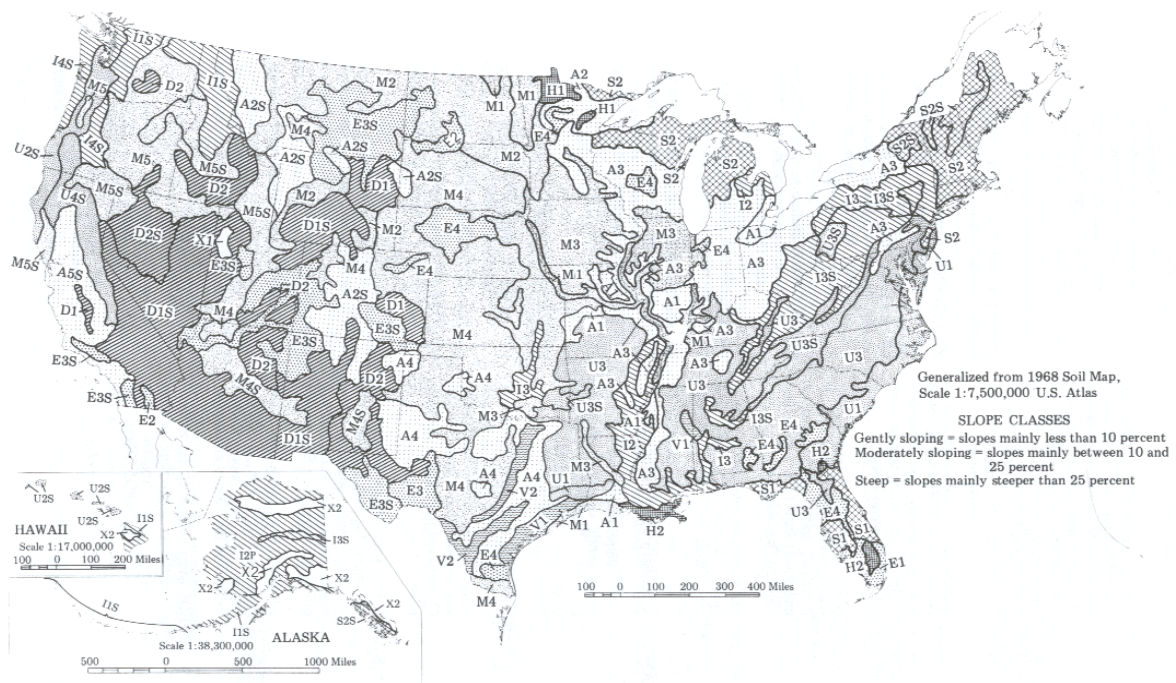


Fig. 2-4 Patterns of soil orders and suborders of the United States. (Courtesy of the U.S. Dept. Agriculture, Soil Conservation Service.)

Chemical Weathering and Soils

Pedogenic controls and regimes

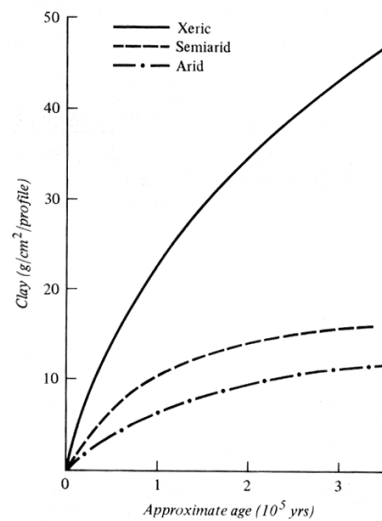
Soil = \int (climate, biota, time, parent material, topography)



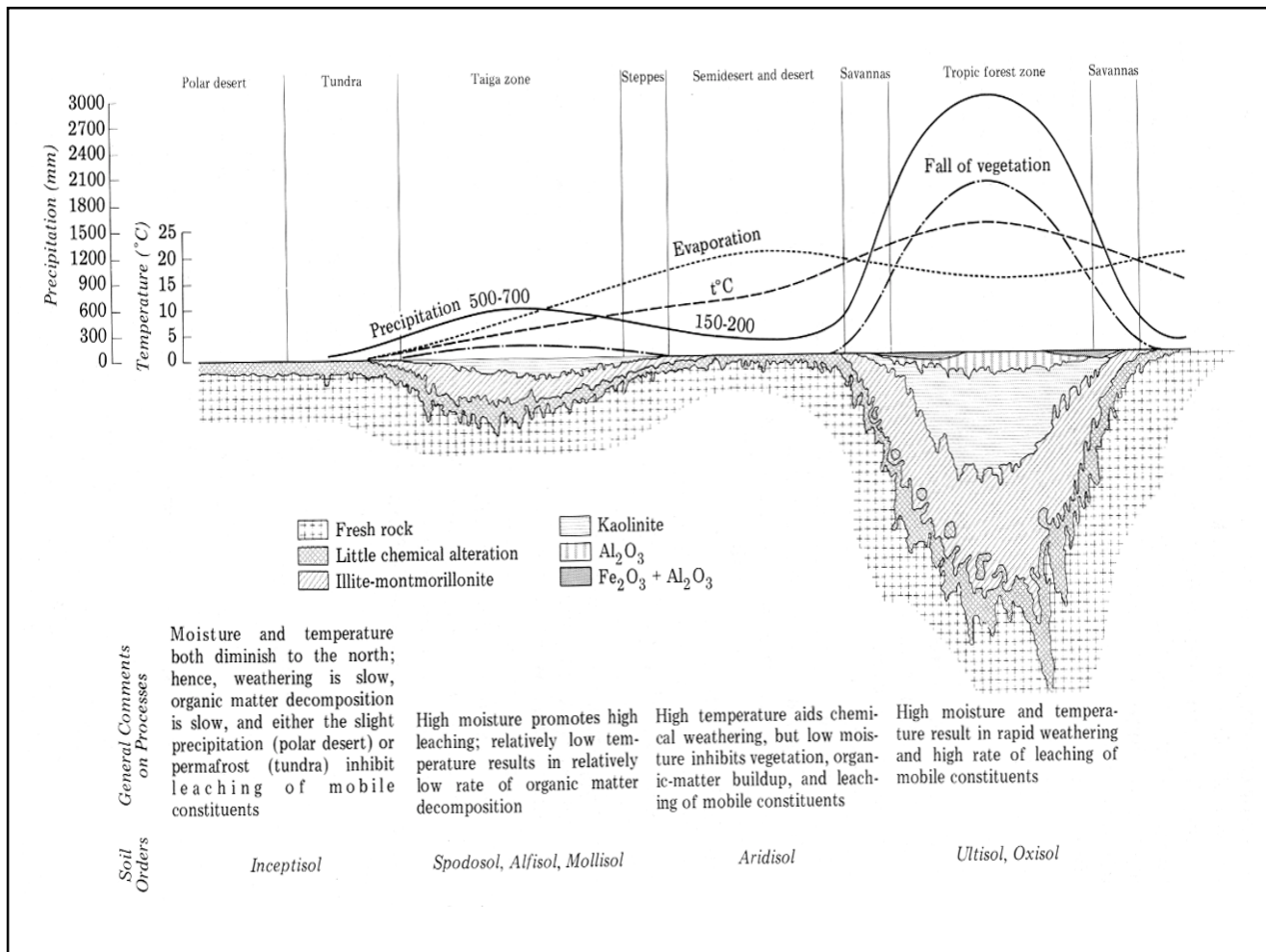
Chemical Weathering and Soils

Pedogenic controls and regimes

Climate



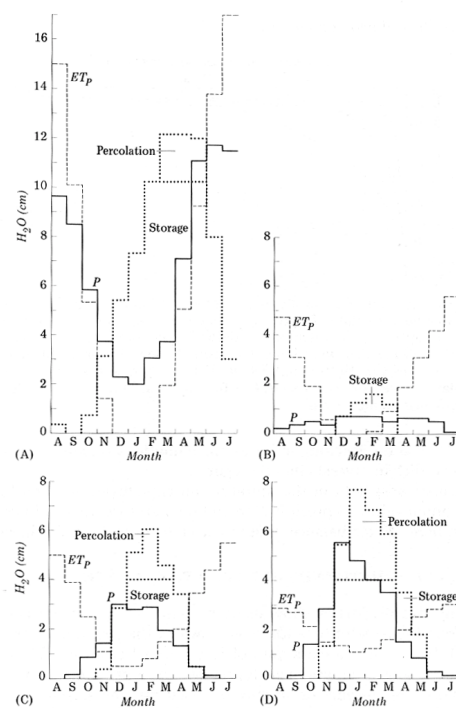
The amount of pedogenic clay as a function of time in three climatic regimes, Southern California. The mean annual precipitations (cm) and temperatures (°C) for the three regimes are, respectively, arid—<12, <22; semi-arid—12–25; 15–22; and xeric—25–75, 12–22. (Taken from McFadden,⁷² Fig. 27.)



Chemical Weathering and Soils

Pedogenic controls and regimes

Climate



Monthly values of precipitation (P), potential evapotranspiration (ET_p), soil-moisture storage in the A and B horizons (10.2-cm available water-holding capacity), and soil-moisture percolation below the Bw or Bt horizon for various climatic stations. Mean temperature as follows:

Station	Mean temperature (°C)	
	Jan.	July
(A) Manhattan, Kan.	-1.6	26.6
(B) Fallon, Nev.	-1.2	22.9
(C) Sacramento, Calif.	7.5	23.2
(D) Half Moon Bay, Calif.	9.9	17.0

Chemical Weathering and Soils

Pedogenic controls and regimes

Parent material

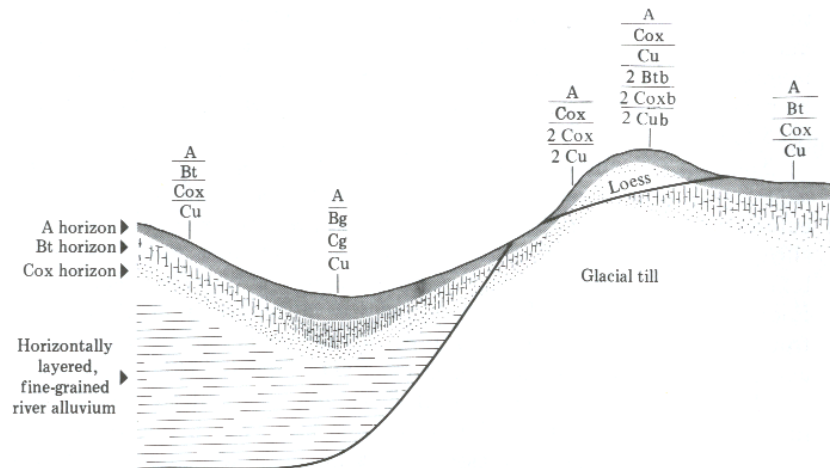


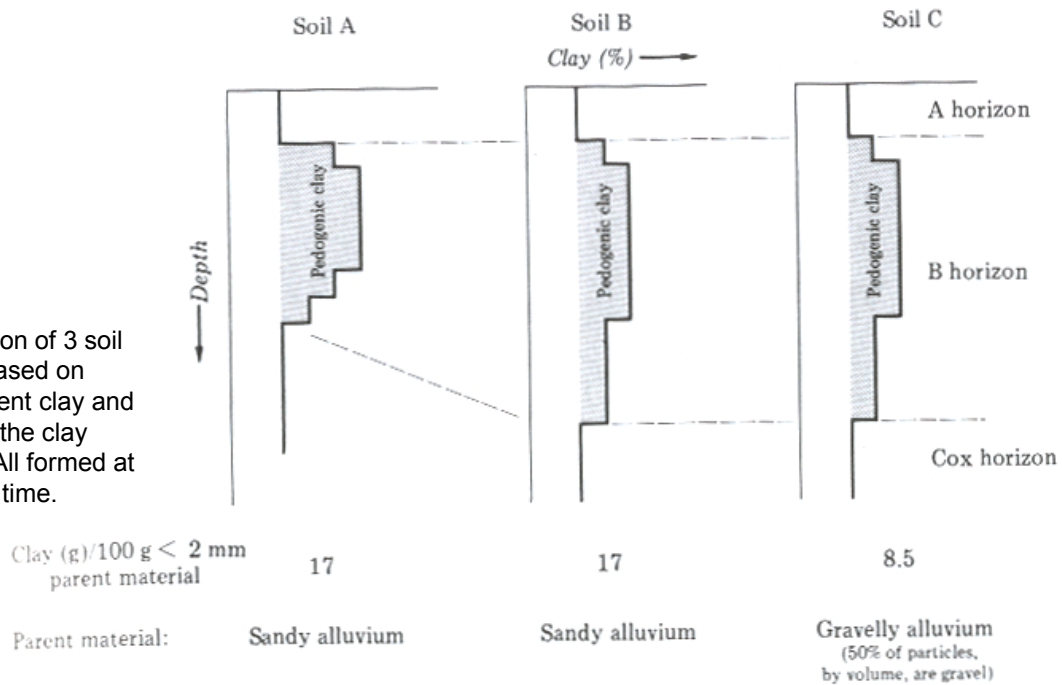
Fig. 1-3 Lateral variation in soil profiles due to lateral variation in environmental conditions and lithology and age of parent materials. Soil horizons parallel the land surface. Those horizons that are developed from river alluvium truncate depositional layering; hence, soil properties are, for the most part, independent of properties of the depositional layers. The soil formed on glacial till was partly truncated by erosion before burial by loess; it grades laterally to the right from a buried to an exhumed soil, because the loess cover has been removed by erosion in that direction.

Chemical Weathering and Soils

Pedogenic controls and regimes

Parent material

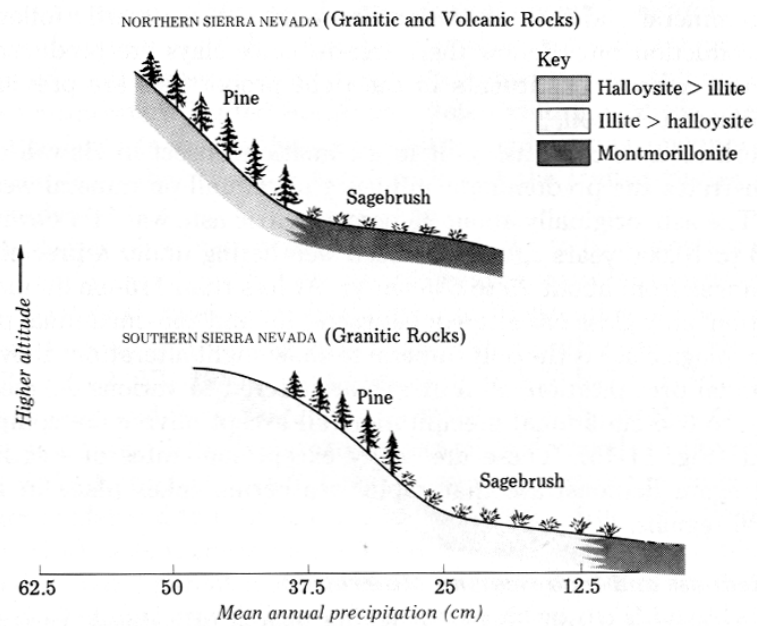
Comparison of 3 soil profiles based on both percent clay and weight of the clay formed. All formed at the same time.



Chemical Weathering and Soils

Pedogenic controls and regimes

Parent material

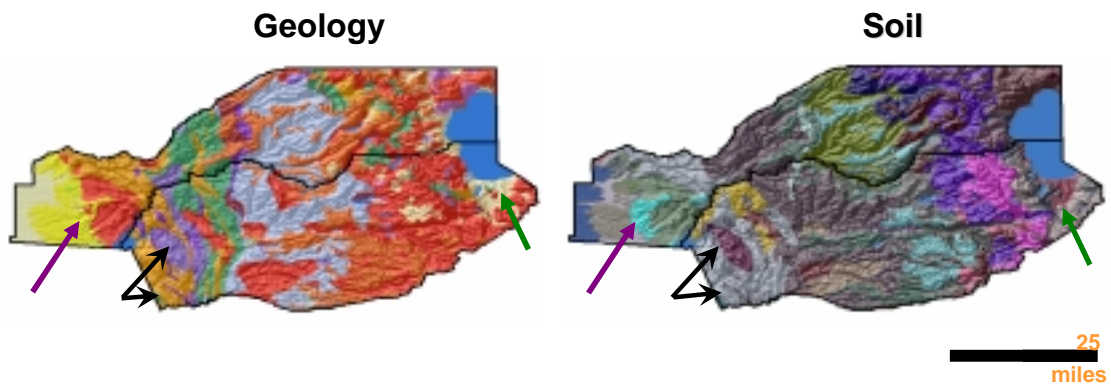


Chemical Weathering and Soils

Pedogenic controls and regimes

Parent material

Correlation of Geology and Soil Distribution for both El Dorado and Placer County

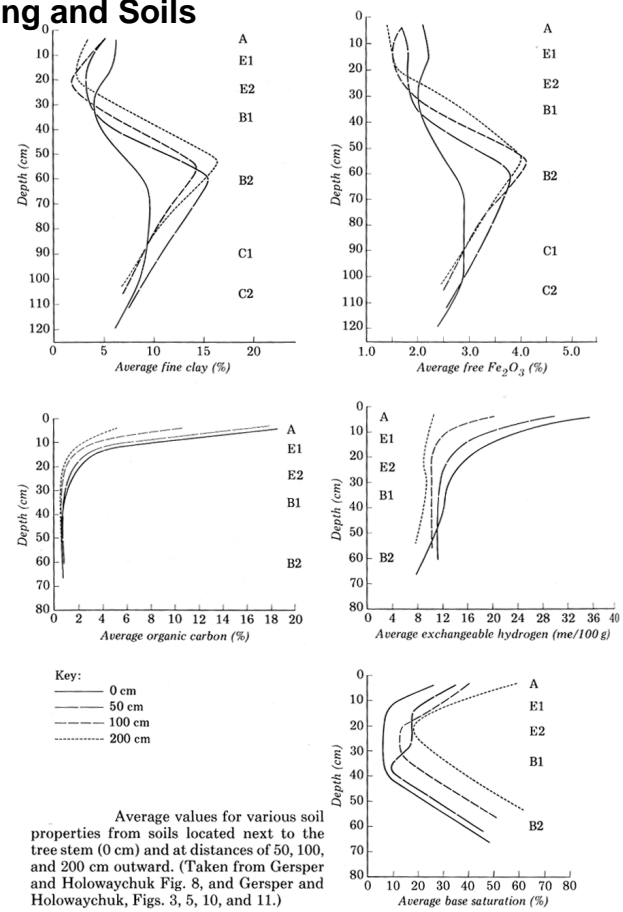


(Similarities shown in same colored arrow)

Chemical Weathering and Soils

Pedogenic controls and regimes

Biotic factor



Chemical Weathering and Soils

Pedogenic controls and regimes

Topography

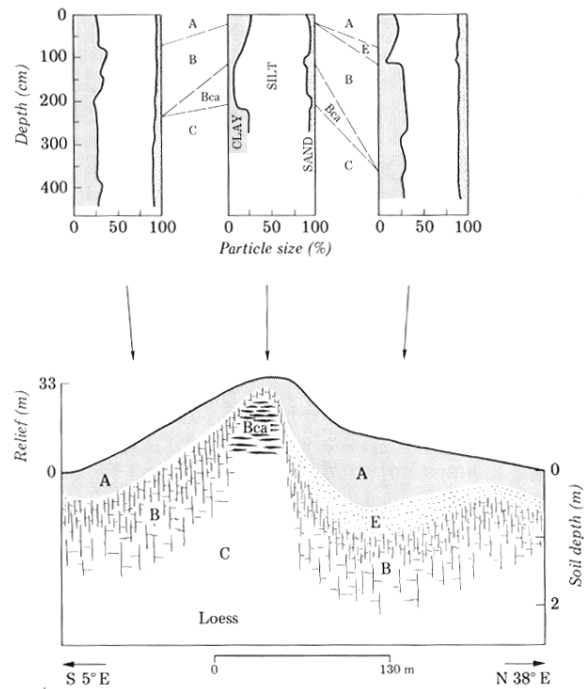


Fig. 9-3 Variation in soil properties with slope orientation in loess hills, eastern Washington. (Taken from Lotspeich and Smith,³⁴ Figs. 2 and 3, © 1953, The Williams & Wilkins Co., Baltimore.)

Chemical Weathering and Soils

Pedogenic controls and regimes

Topography

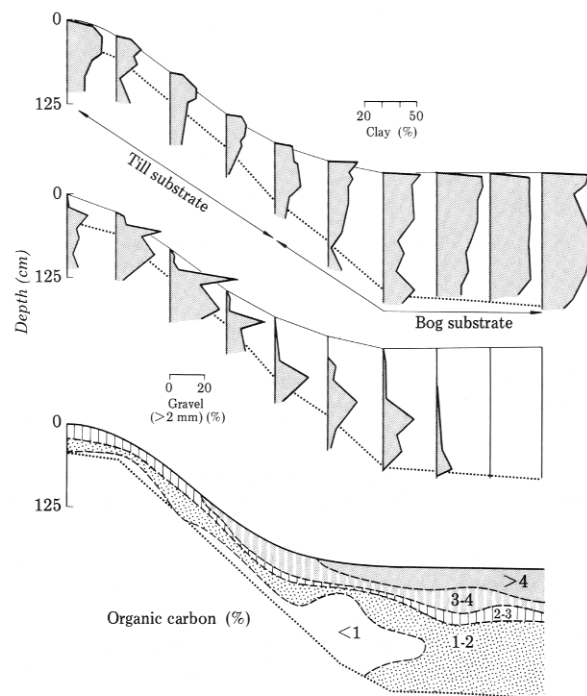
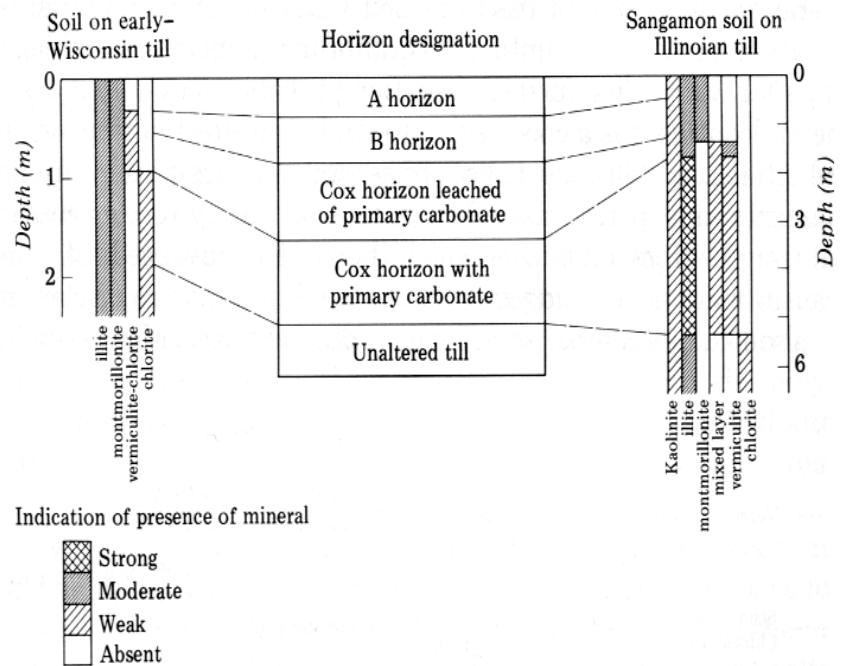


Fig. 9-7 Relationship of particle size and organic carbon with topographic position in a closed depression and on surrounding hillslopes, Iowa. (Taken from Walker,⁵⁷ Figs. 19, 30, and 31.) The dotted line separates hillslope sediments from till on the hillslopes, and younger from older deposits in the low-lying area.

Chemical Weathering and Soils

Pedogenic controls and regimes

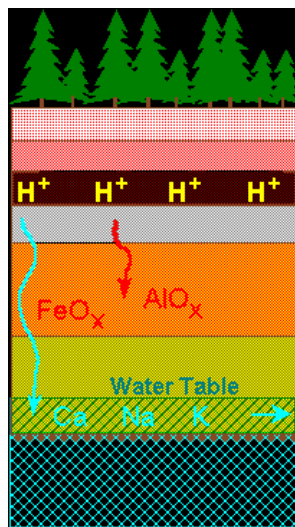
Time



Chemical Weathering and Soils

Pedogenic controls and regimes

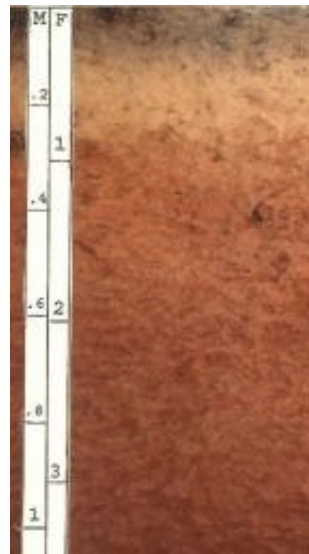
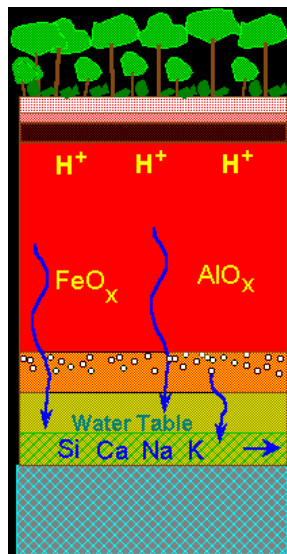
Podzolization - downward translocation of iron, aluminum, and organic matter to form a eluvial E horizon overlying a tan illuvial B horizon.



Chemical Weathering and Soils

Pedogenic controls and regimes

Laterization - extreme chemical alteration of the original parent material.



Chemical Weathering and Soils

Pedogenic controls and regimes

Calcification - a general buildup of calcium carbonate in the soil profile.



Chemical Weathering and Soils

Pedogenic controls and regimes

Salinization - soluble salts precipitate from water and accumulate in the soil.
Saline soils are common in desert and steppe climates.



Chemical Weathering and Soils

Geomorphic significance of soils

They are used primarily

- in the subdivision of a local succession of deposits,



Chemical Weathering and Soils

Geomorphic significance of soils

They are used primarily

- to provide data on the lengths of time that separate periods of deposition,
- to facilitate short- and long-term correlation, and, of recent discovery,
- to date Quaternary events (faulting, climate change, etc.)
- archeological studies
- Reconstruct climates



Chemical Weathering and Soils

Geomorphic significance of soils

Soils that form on a landscape of the past are *paleosols*.

They can be of three types:

1. *buried soils* are developed on a former landscape and subsequently covered by younger alluvium or rock;
2. *relict soils* were not subsequently buried but still exist at the surface; and
3. *exhumed soils* were at one time buried but have been re-exposed when their cover was stripped by erosion.

