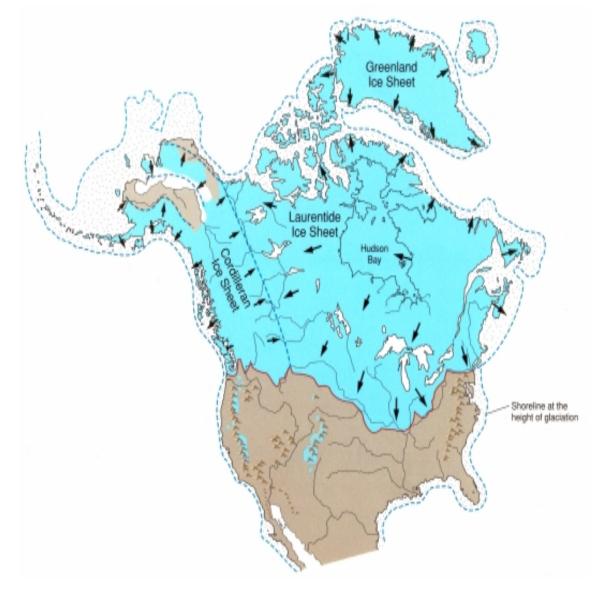
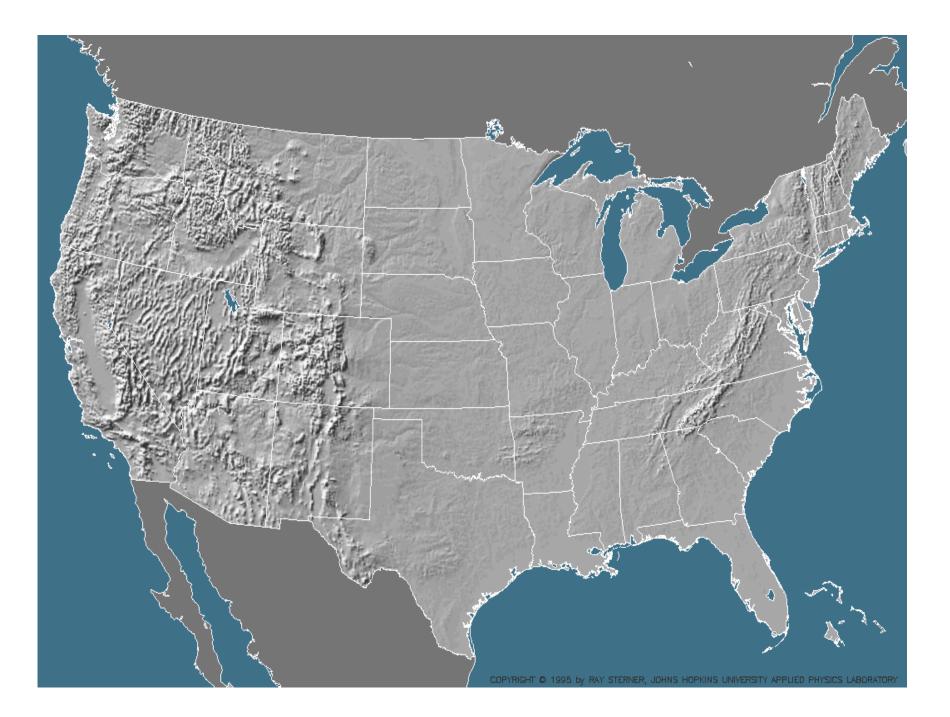
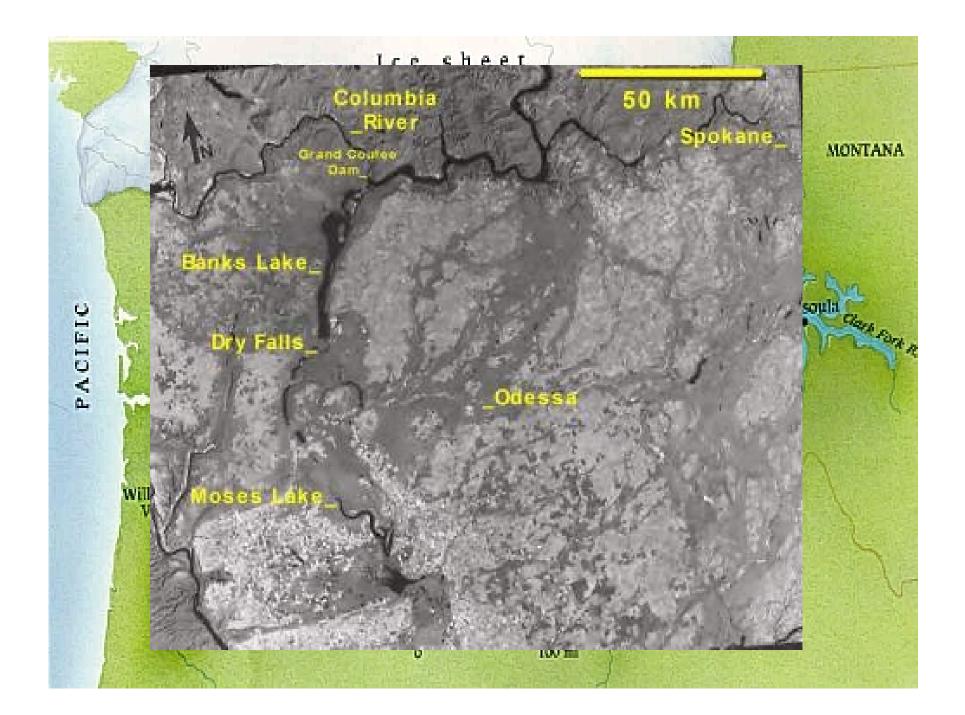
Glaciers

Immense and extensive glaciers covered as much as a 1/3 of the Earth's land surface







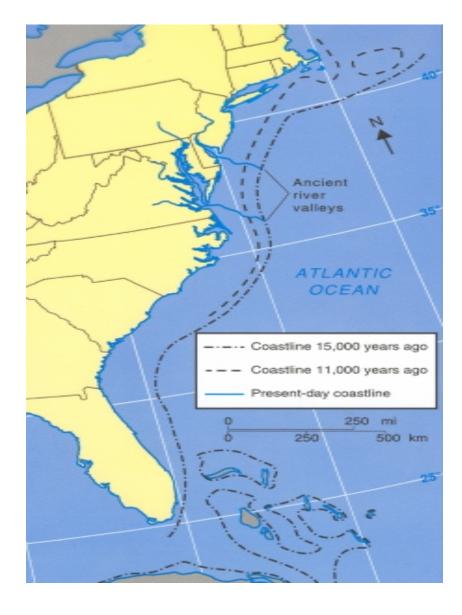


Glaciers



Glaciers

World-wide climatic changes during the colder glacial ages distinctly altered landscapes in areas far from the glacial boundaries.



Glaciers

This lowering of world-wide sealevels likely resulted in the original peopling of North America.



Glaciers

Glacier – a mass of ice that moves over land under its own weight due to gravity.

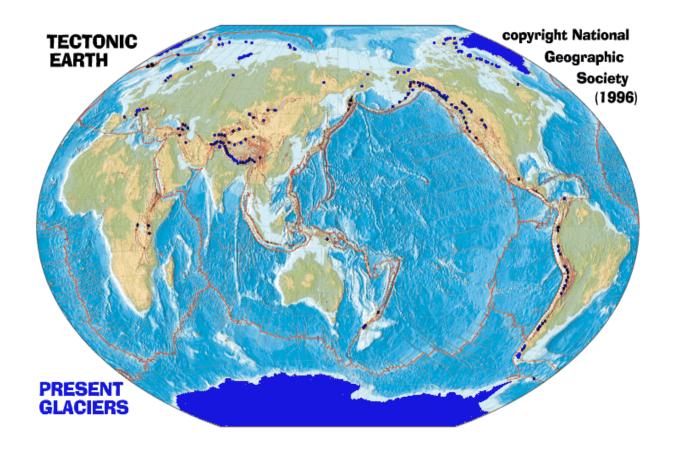


Simply put:

Snow accumulation > snow melting

Glaciers – global distribution

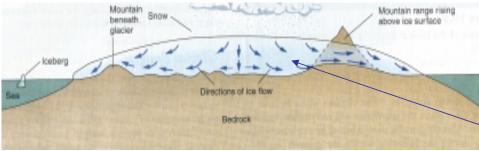
Glaciers are found in both polar regions, where there is little melting during the summer, and in temperate regions that have heavy snowfall during the winter months.



Glaciers – types

Two primary types of glaciers, *valley glaciers* are confined to valleys and owe their shape and size to the confines of the valleys they occupy.





Glaciers that are not confined to the landscape are called *ice-sheets*.

Glaciers – classification

Morphological classification:

- Cirque glaciers
- Valley glaciers
- Ice sheets
 - Mountain ice sheets
 - Piedmont glaciers

Dynamic classification:

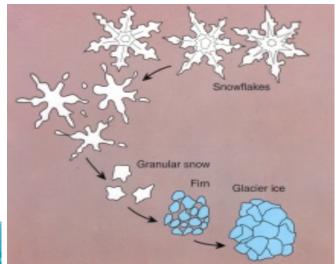
- Active glaciers
- Passive glaciers
- Dead glaciers

Thermal classification:

- •Temperate glaciers
- •Polar glaciers
- •Sub-polar glaciers
- •High-polar glaciers

Glaciers – dynamics

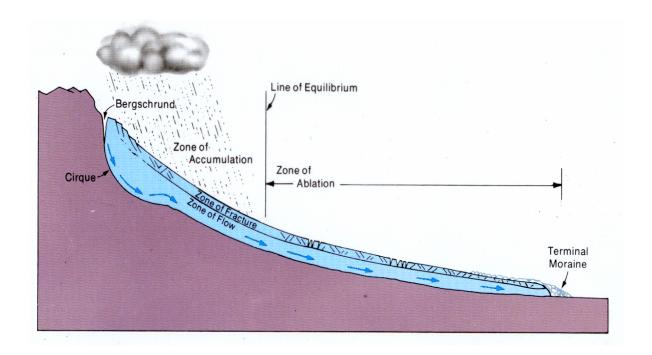
Glacial ice develops by the *compression and compaction of snow and ice*.





Glaciers – dynamics

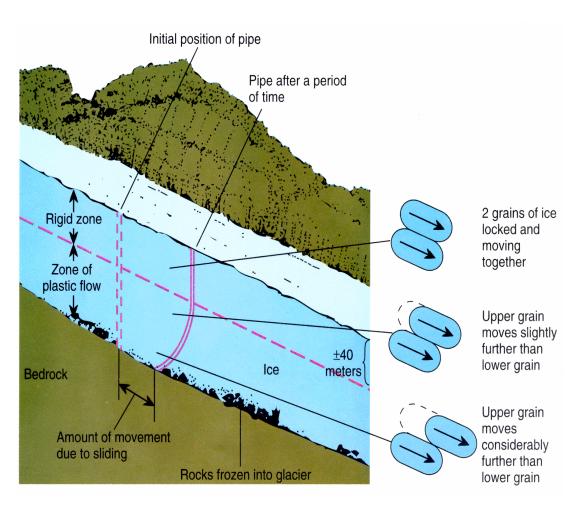
A balance exists between the **zone of accumulation** and the **zone of wastage** on a glacier. The line separating these two conditions is called the <u>line of equilibrium</u>.



Glaciers – dynamics

Glaciers move by shearing within the ice mass.

A tremendous amount of sediment is produced as the ice moves over rock and the debris on the floor of the valley.



Glaciers – movement of glaciers

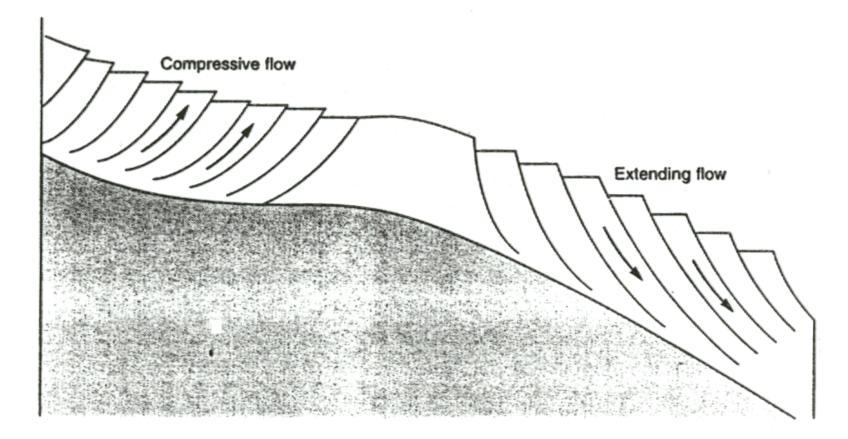
Glaciers move by one of two processes:

Internal deformation of the ice called *creep*, and

sliding of the glacier along its base and sides.

A *surging glacier* is one in which sudden, brief, large scale ice displacements periodically occur.

Glaciers – movement of glaciers - Extending and Compressive Flow



Glaciers – ice structures

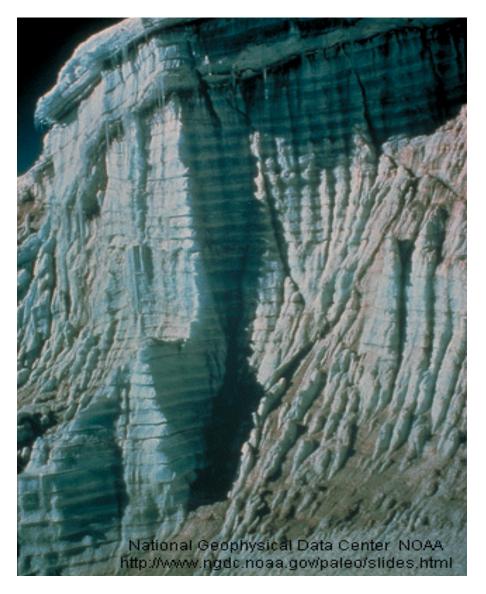
Glaciers usually develop a variety of structures that *develop during growth of the glacial mass*, which we can call primary structures.



Glaciers – ice structures

Stratification

Primary structures in glaciers appear as discernible layers or bands within the ice. The layering results from processes that *reflect an annual cycle of snow accumulation and ablation above the firn line*.



Glaciers – ice structures

Foliation

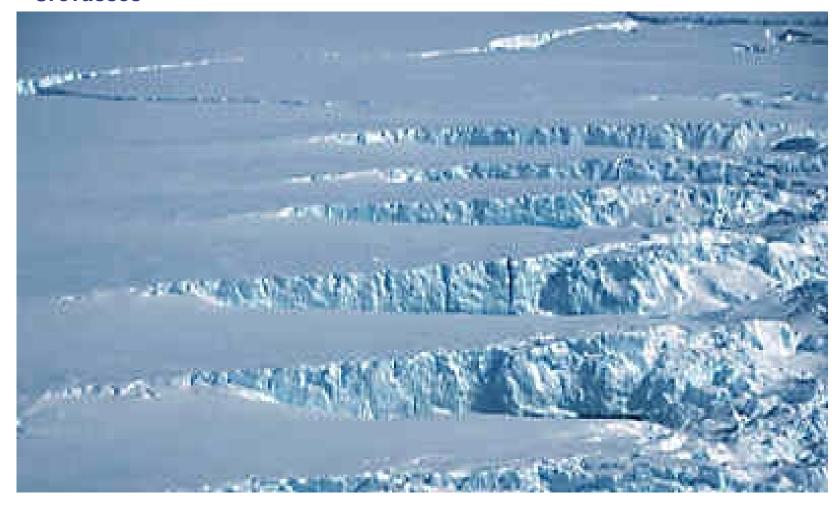
Is produced by shearing during ice motion. It is sometimes difficult to distinguish from primary stratification because both types display similar grain size or textures.



Glaciers – ice structures Crevasses

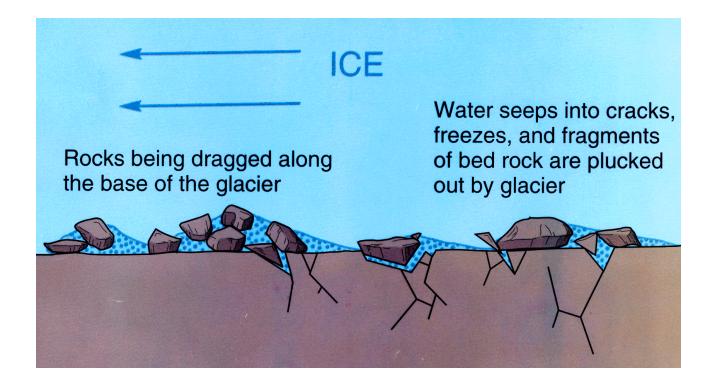


Glaciers – ice structures Crevasses



Glaciers – erosional processes

Glacial erosion is accomplished primarily by two processes, a scraping action called *abrasion* and a dislodgment or lifting action called *quarrying* or *plucking*.



Glaciers – erosional processes

A *cirque* by any name is still a deep erosional recess with steep and shattered walls that is usually located at the head of a mountain valley.



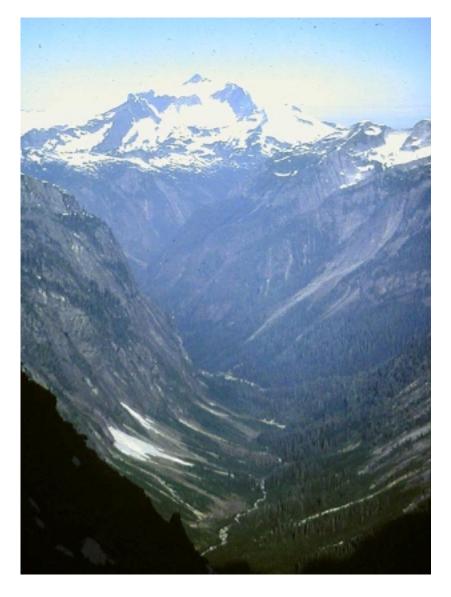
Glaciers – erosional processes

The bowl shaped basin often contain lakes, called *tarns*.



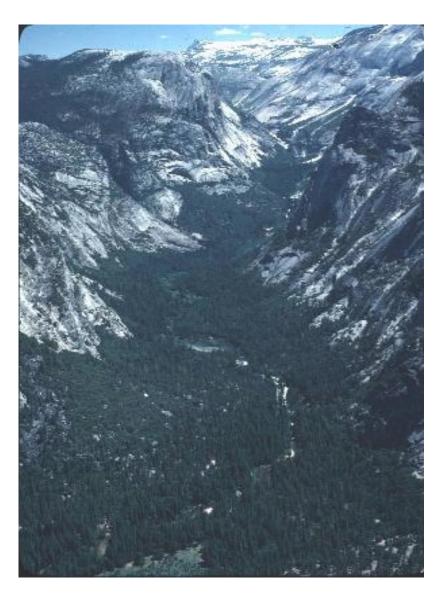
Glaciers – erosional processes

U-shaped glacial valley



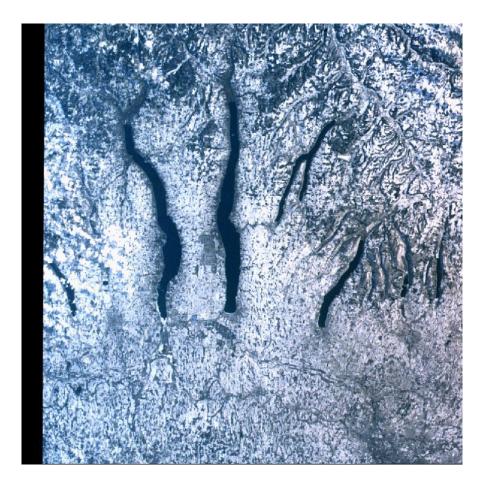
Glaciers – erosional processes

U-shaped glacial valley

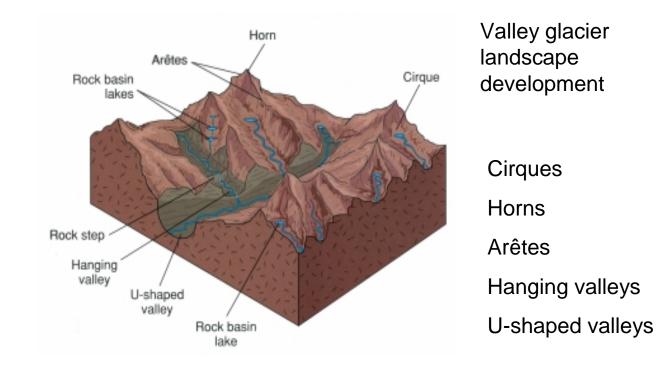


Glaciers – erosional processes

Glaciated valleys and troughs that contain lakes are sometimes called *paternoster or finger lakes*



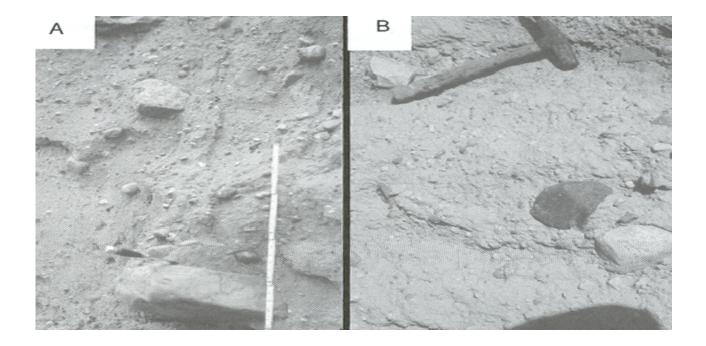
Glaciers – erosional landscapes



Glaciers – depositional processes and landscapes

Nonstratified drift

Sediments originating directly from glacial ice characteristically has no discernible stratification and is generally called *till*.



Μ

Glaciers – depositional processes and landscapes

Nonstratified drift

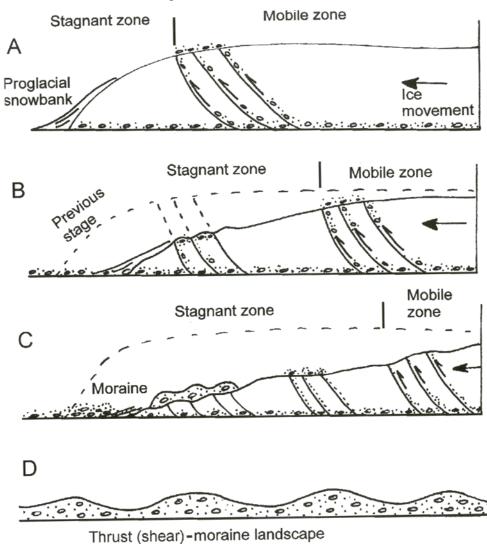
Each transport sub-environment produces till with different characteristics. *Supra-glacial* till has a texture dominated by coarse, angular clasts

Sub-glacial till is more compact and contains a higher percentage of finegrained sediments.

En-glacial load can be deposited directly on the bedrock floor as sub-glacial matter, and a considerable thickness of the glacier will remain above it at the moment of deposition.

Glaciers – depositional processes and landscapesStratified DriftStagnant zone

Moraines



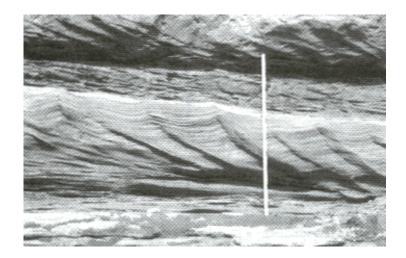
Glaciers – depositional processes and landscapes

Stratified Drift

Fluvioglacial because running water is involved in its origin even though the water may not always be confined in discrete channels.

Fluvioglacial deposits are also distinguished from till in that they are usually sorted and the clasts contained in the mass are more rounded. Sorting is also partly a function of:

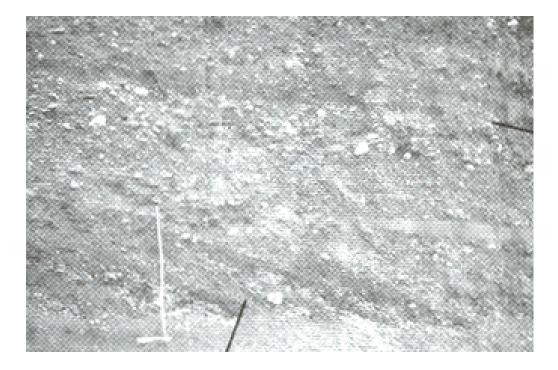
- the energy possessed by the meltwater,
- the distance of transport, and
- the continuity of the sorting process.



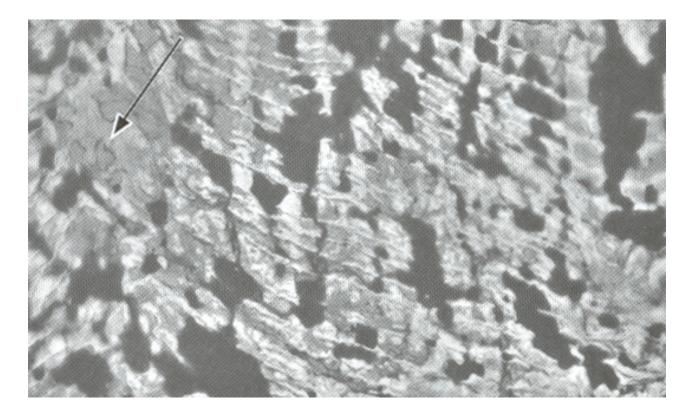
Glaciers – depositional processes and landscapes

Stratified Drift

Sediments deposited beyond the terminal margin of the ice are formed in the proglacial environment and are often referred to as *outwash*.



Glaciers – depositional processes and landscapes Stratified Drift end (recessional) moraines and terminal moraine.

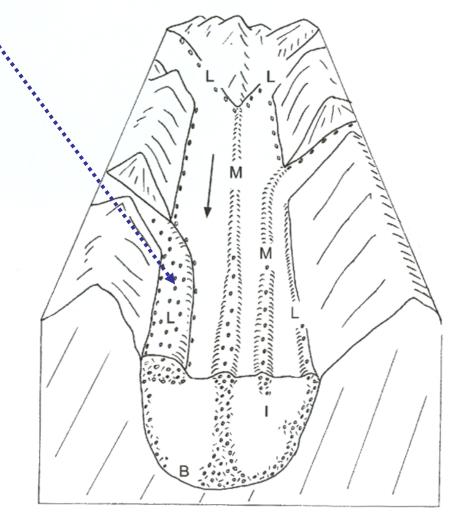


Central Canada

Glaciers – depositional processes and landscapes

Stratified Drift

In valley glacier systems, *lateral*, *moraines* occur on both sides of the valley.

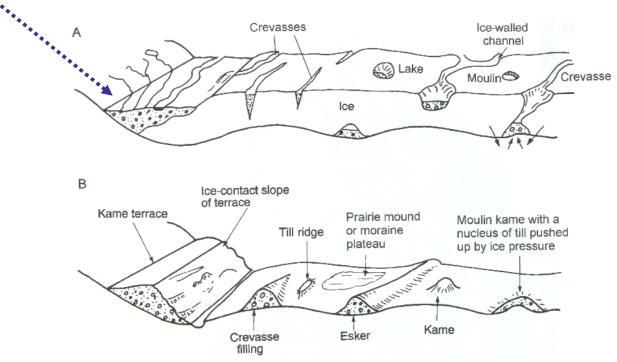


Glaciers – depositional processes and landscapes



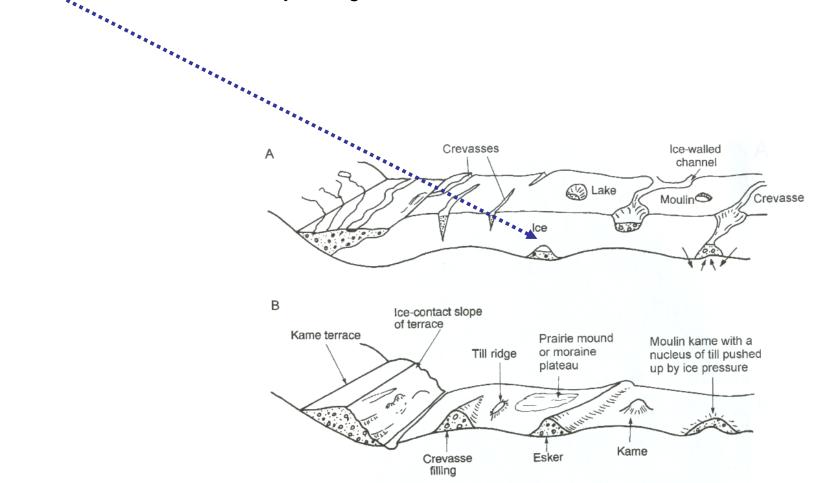
Glaciers – depositional processes and landscapes

Kames are only one of many forms with essentially the same origin. *Kame terraces* originate from drift deposited in narrow lakes or stream channels between the valley side and the lateral edge of the stagnating ice. When the supportive ice disappears, the inner edge of the deposit collapses into the terrace scarp.

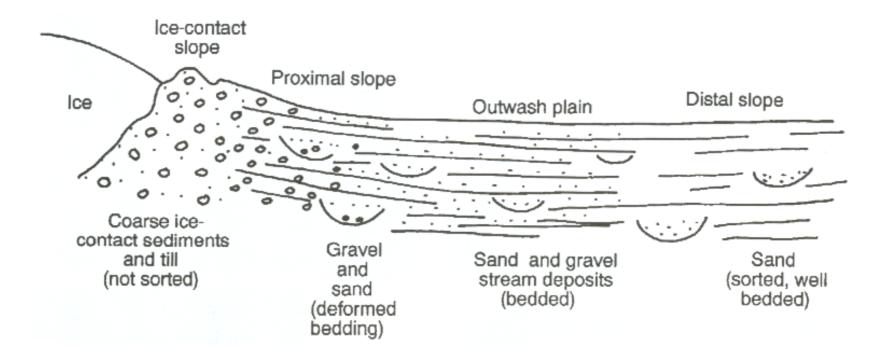


Glaciers – depositional processes and landscapes

Eskers describe a wide variety of ridged ice-contact features.



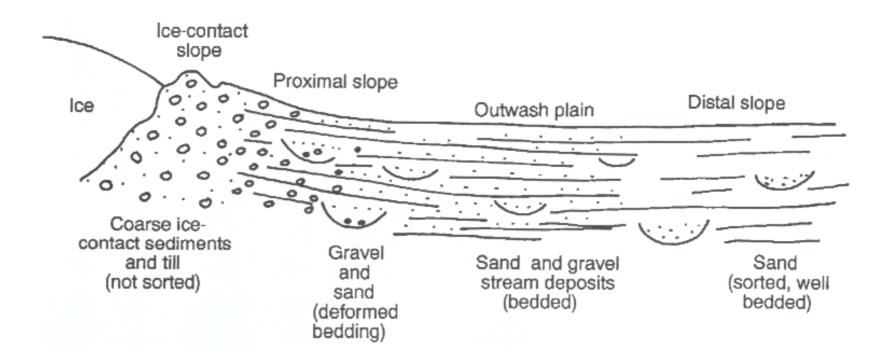
Glaciers – depositional processes and landscapes outwash plain



Glaciers – depositional processes and landscapes

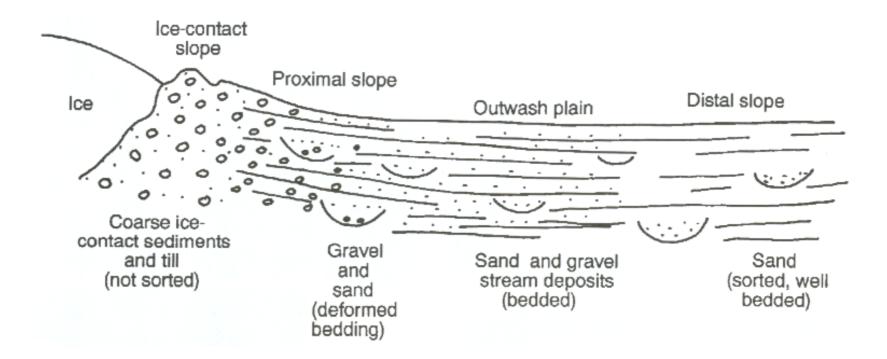
Three distinct zones recognized on the outwash plain.

The *proximal zone,* closest to the ice, is usually transversed by only a few main rivers that flow in well-defined entrenched channels.



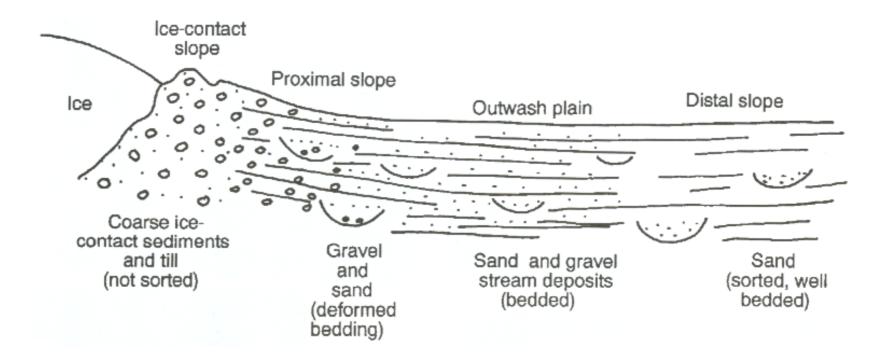
Glaciers – depositional processes and landscapes

In the *intermediate zone*, the channels become wide and shallow and distinctly braided, and the entire depositional network shifts its position rapidly from side to side.



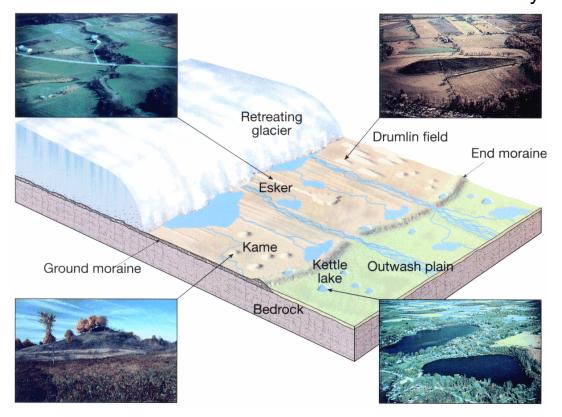
Glaciers – depositional processes and landscapes

Downstream the system changes gradually into the *distal zone*, where channels become so shallow that the rivers may merge into a single sheet of water during high flow.

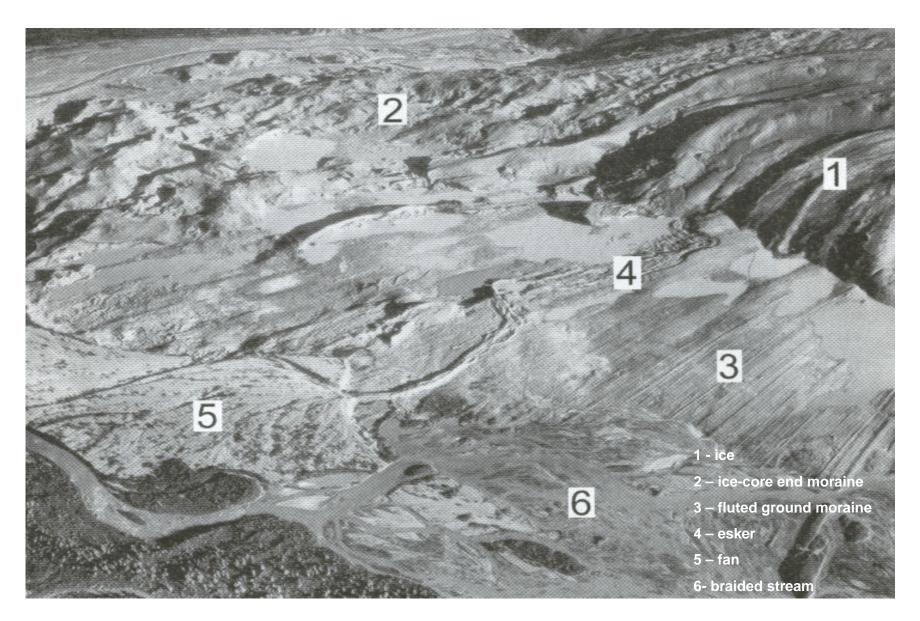


Glaciers – depositional landscapes

Glaciers are also depositional machines. They deposit all of that sediment that they erode. Although all glaciers both erode and deposit sediment, the ice-sheets drop substantial loads of sediment as they melt. Produces many of today's landscapes.

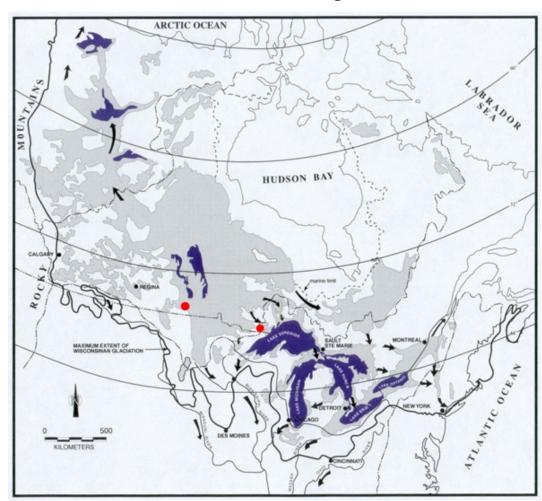


Moraine Kettle lakes Eskers Drumlins

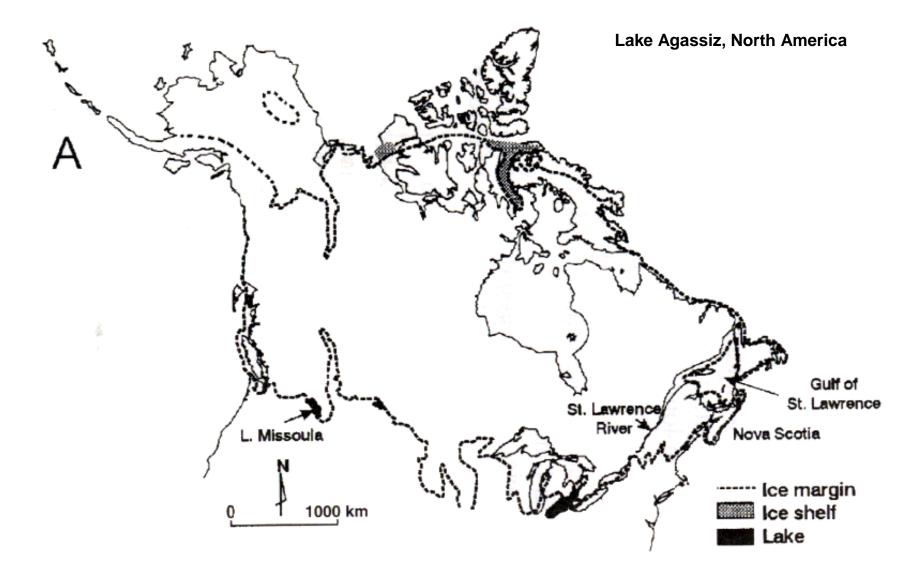


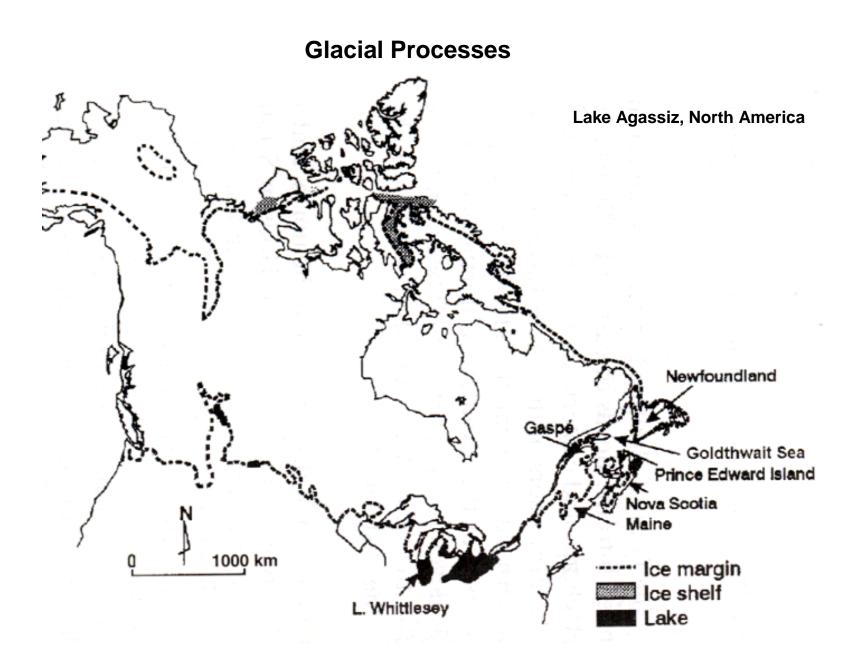
Glaciers – pro-glacial lakes

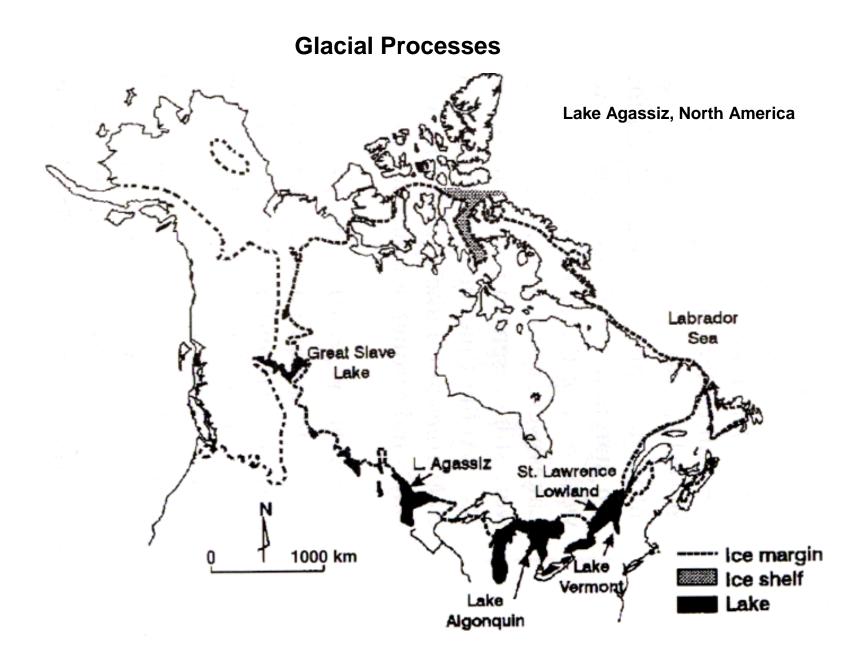
As glacial ice masses melt, they release a tremendous amount of water that often collects at the melting edge of the glacier due to isostatic depression of the land mass. These areas form *pro-glacial lakes* and can cover very large tracts of land.

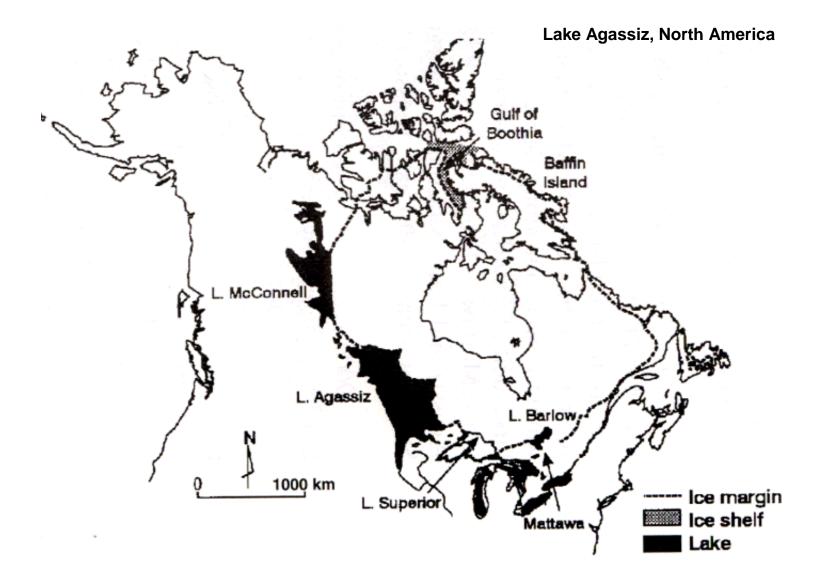


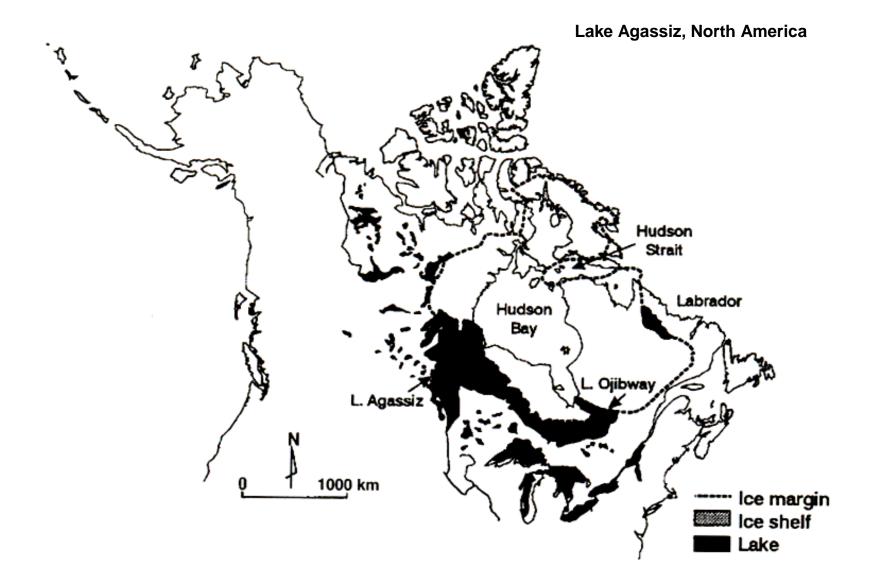
Lake Agassiz, North America

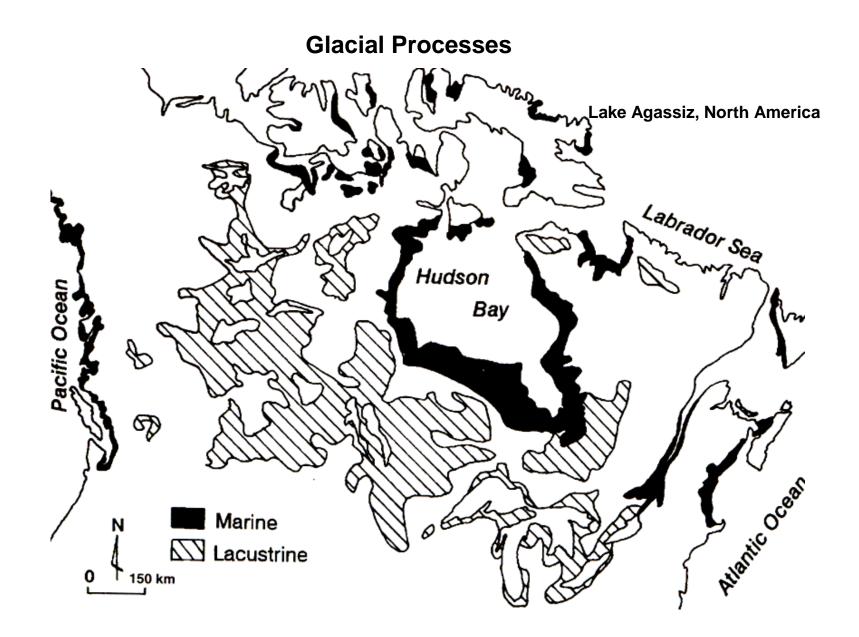


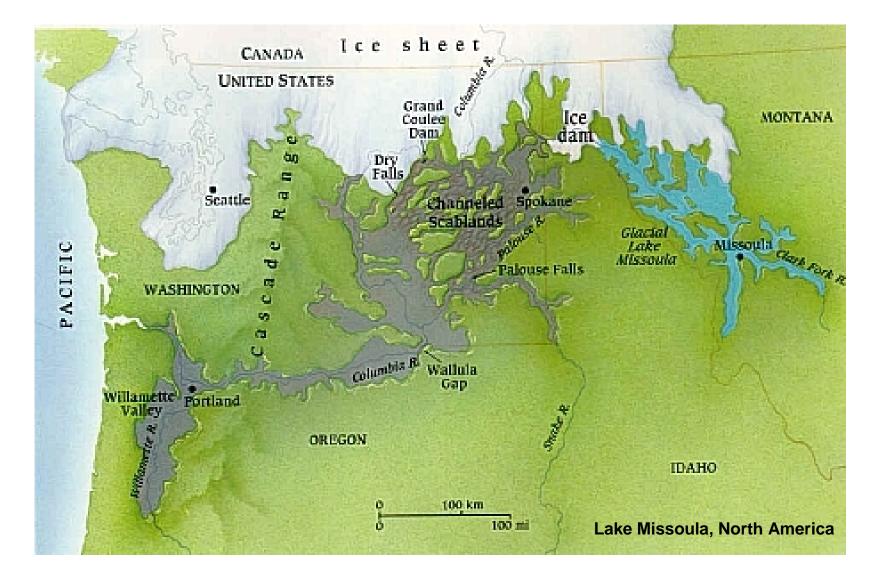












Glacial Processes

Lake Missoula, North America



Lake Missoula, North America



Strandlines (ancient lakeshores) on Mt. Jumbo/ NPS photo

GLACIERS

Why does the Earth experience glacial epics at all?

So there were vast ice-sheets that covered substantial portions of the Earth's landmass in the geologic past. Obviously it must be too warm today to support such ice sheets because we don't see them, so what causes the Earth to get cool enough to start growing glaciers?

Possible cause:

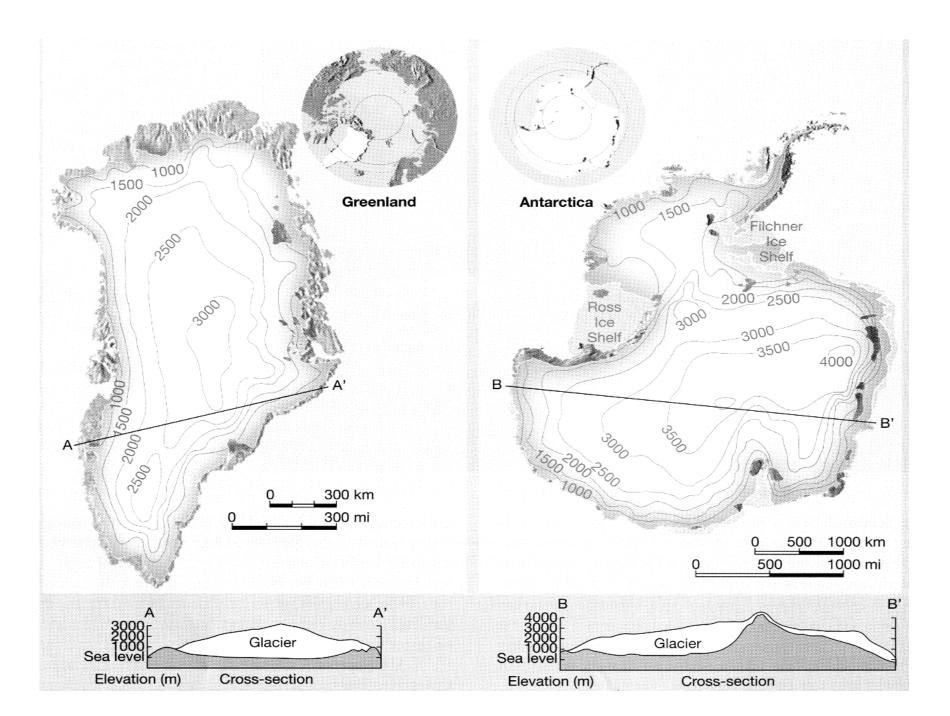
Changes in the position of the continents

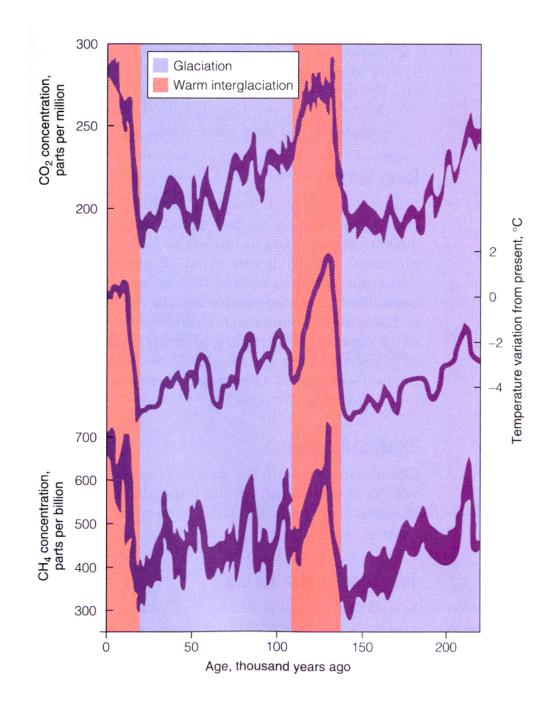
Changes in the circulation of sea water

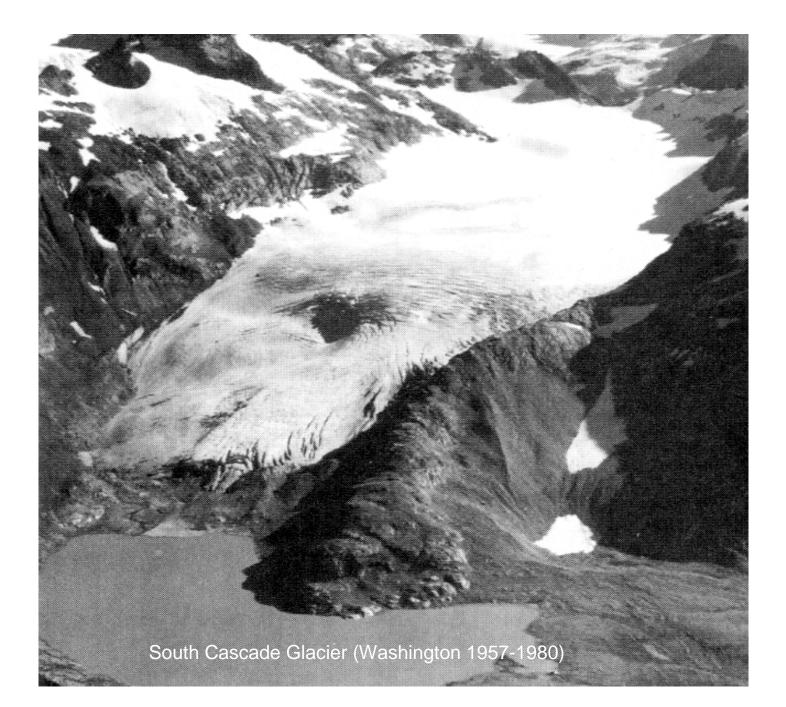
Milankovitch cycles

Changes in the atmosphere – reduction in greenhouse gases

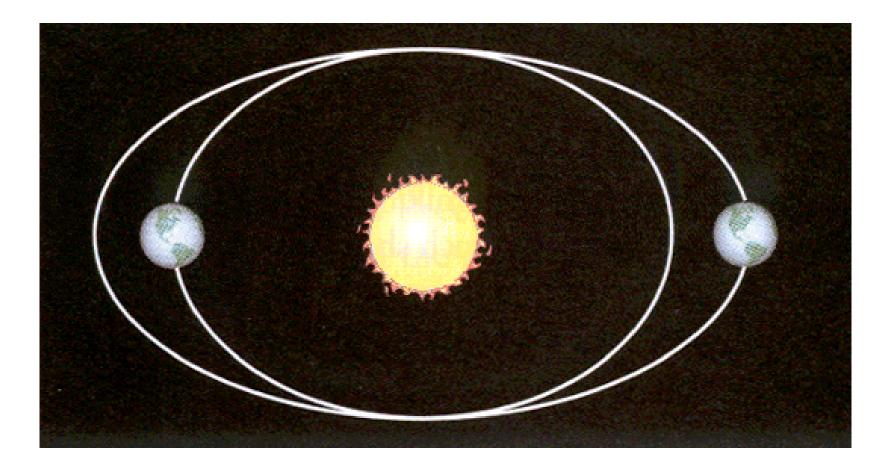
links



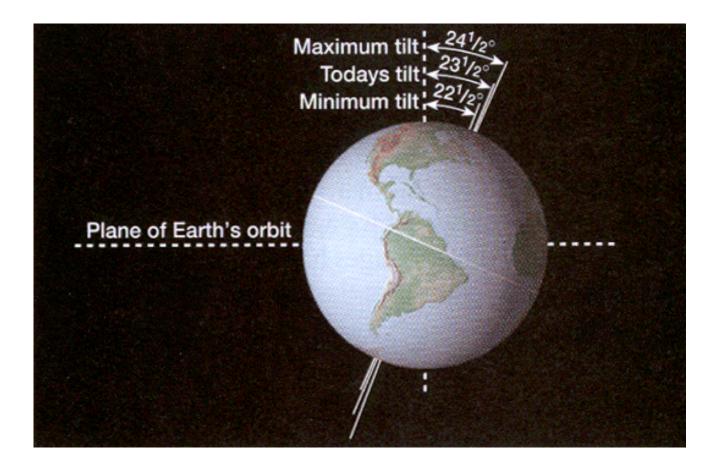




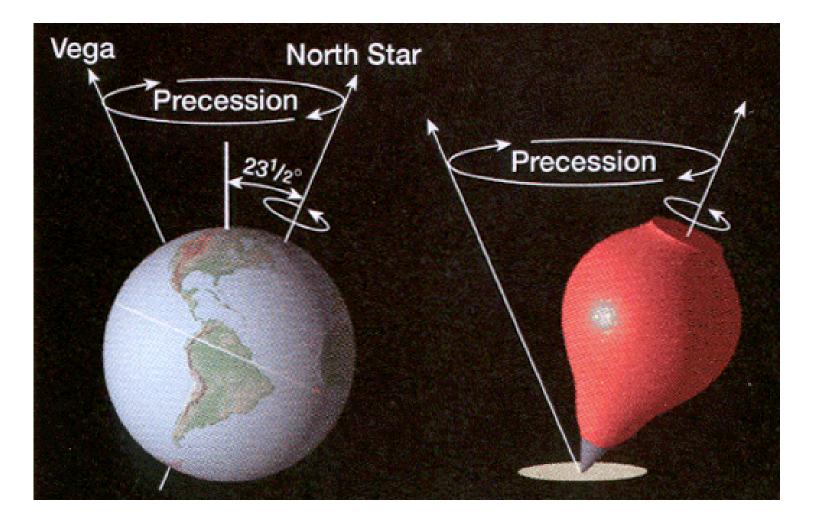
Eccentricity of the Earth's orbit – 96,000 years



Axial Tilt – 42,000 years



Precession of the equinoxes – 21,000 years



Periodicity

